



How Well do we Understand the Causes of Drought?

**Attribution Workshop
17-18 August 2010
Broomfield, CO**

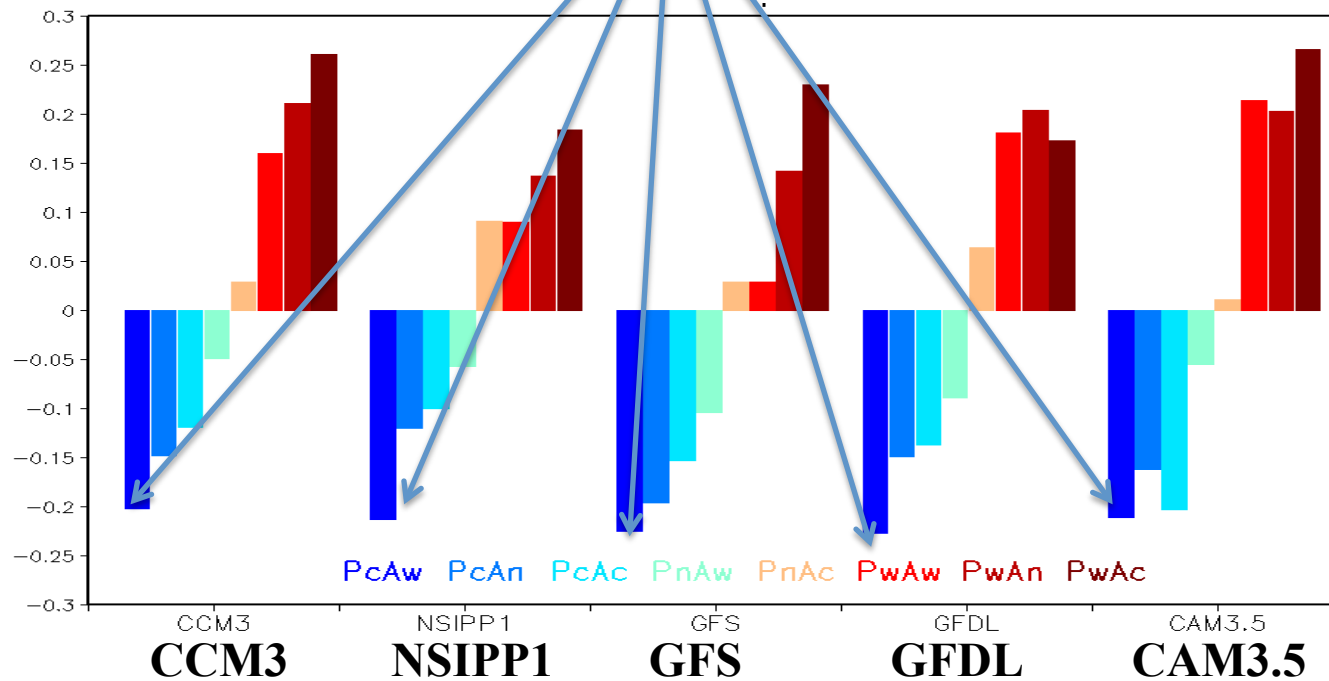
Siegfried Schubert¹, Hailan Wang^{1,2}, Max Suarez¹, and Randal Koster¹

¹NASA/GMAO; ²UMBC/GEST

Some Indication of Our Understanding

- A cold tropical Pacific combined with a warm tropical Atlantic should lead to drought in the Great Plains

Great Plains Precip Anomaly: AGCM results with Idealized SST Anomalies*



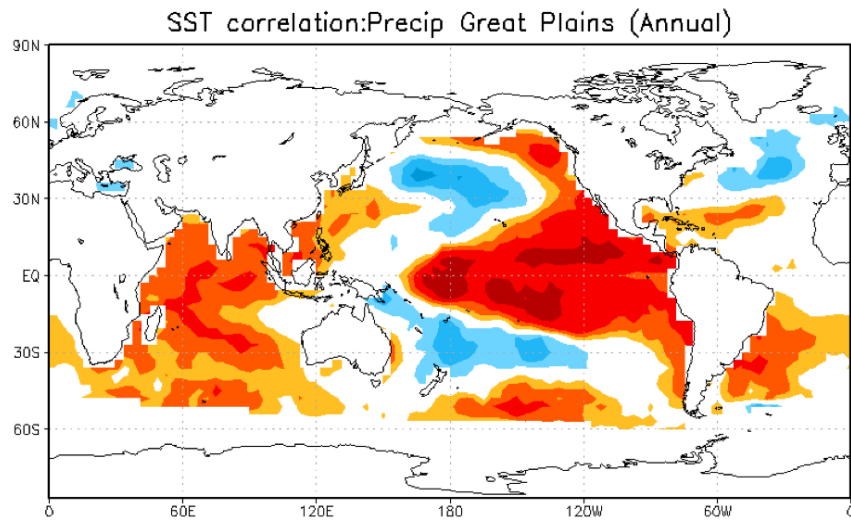
- **This summer: cold Pacific, warm Atlantic and wet in Great Plains!**

Outline

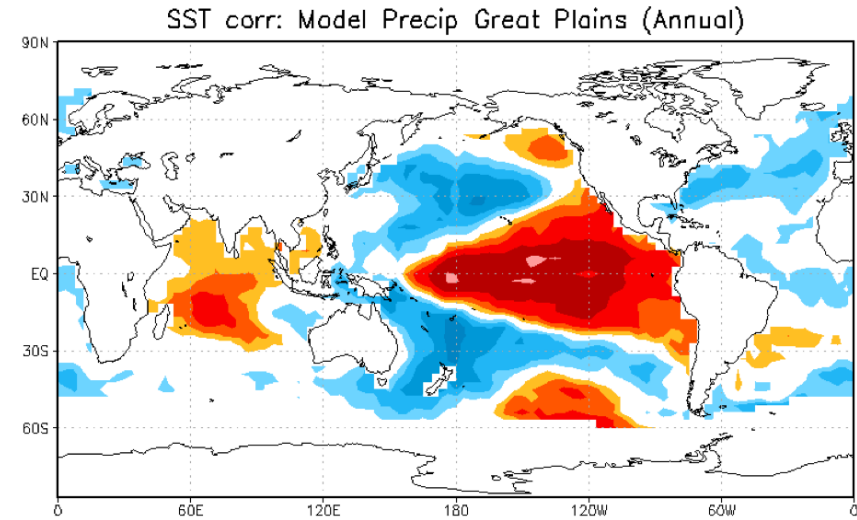
- Decadal drought
 - Importance of transition seasons
- Separating decadal drought from trends
 - Separating decadal SST variability from SST trends
- Sub-seasonal drought/heat waves
 - Important role of stationary Rossby waves

Correlations Between Great Plains precipitation and annual mean SST

Observations



NSIPP-1 AGCM

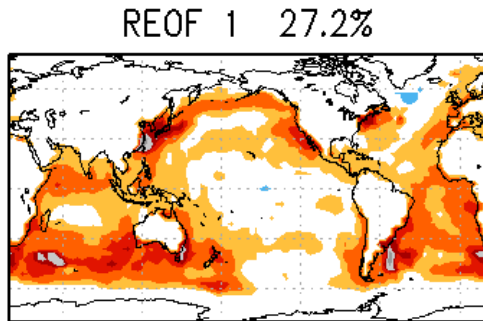


Correlations between annual mean observed (left panel, 1901-2004) and simulated (right panel, 1902-2004) Great Plains precipitation and annual mean SST. For the AGCM simulations, the values are the average of 14 correlations produced for each ensemble member.

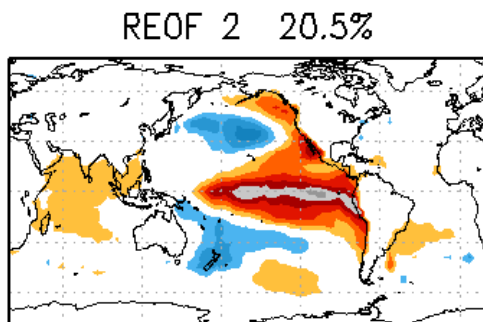
Decadal drought

Leading EOFs and Time series (annual mean SST - 1901-2004)

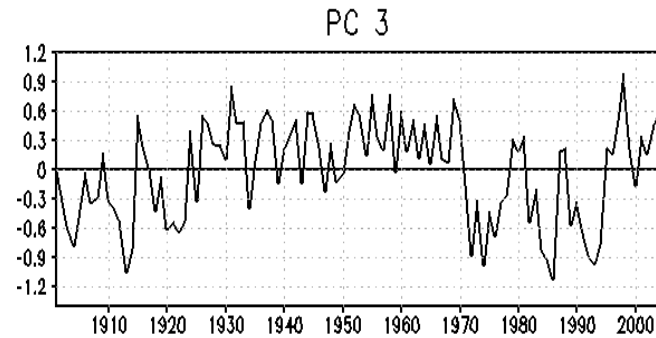
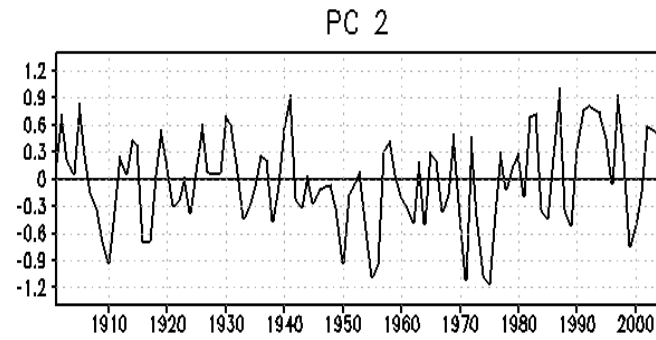
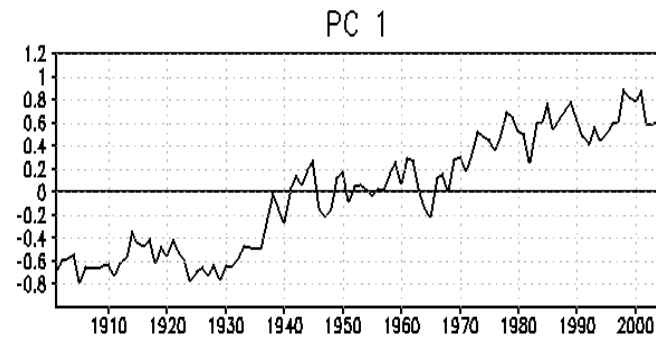
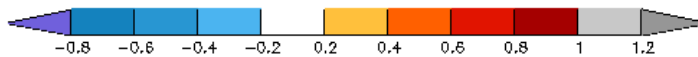
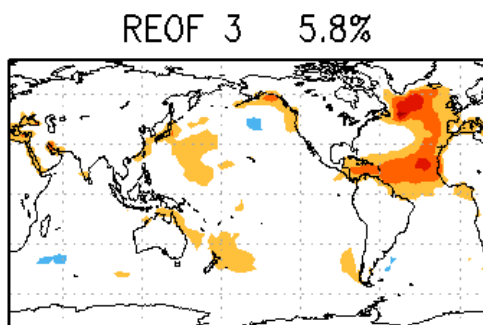
Linear
Trend
Pattern
(LT)



Pacific
Pattern
(Pac)

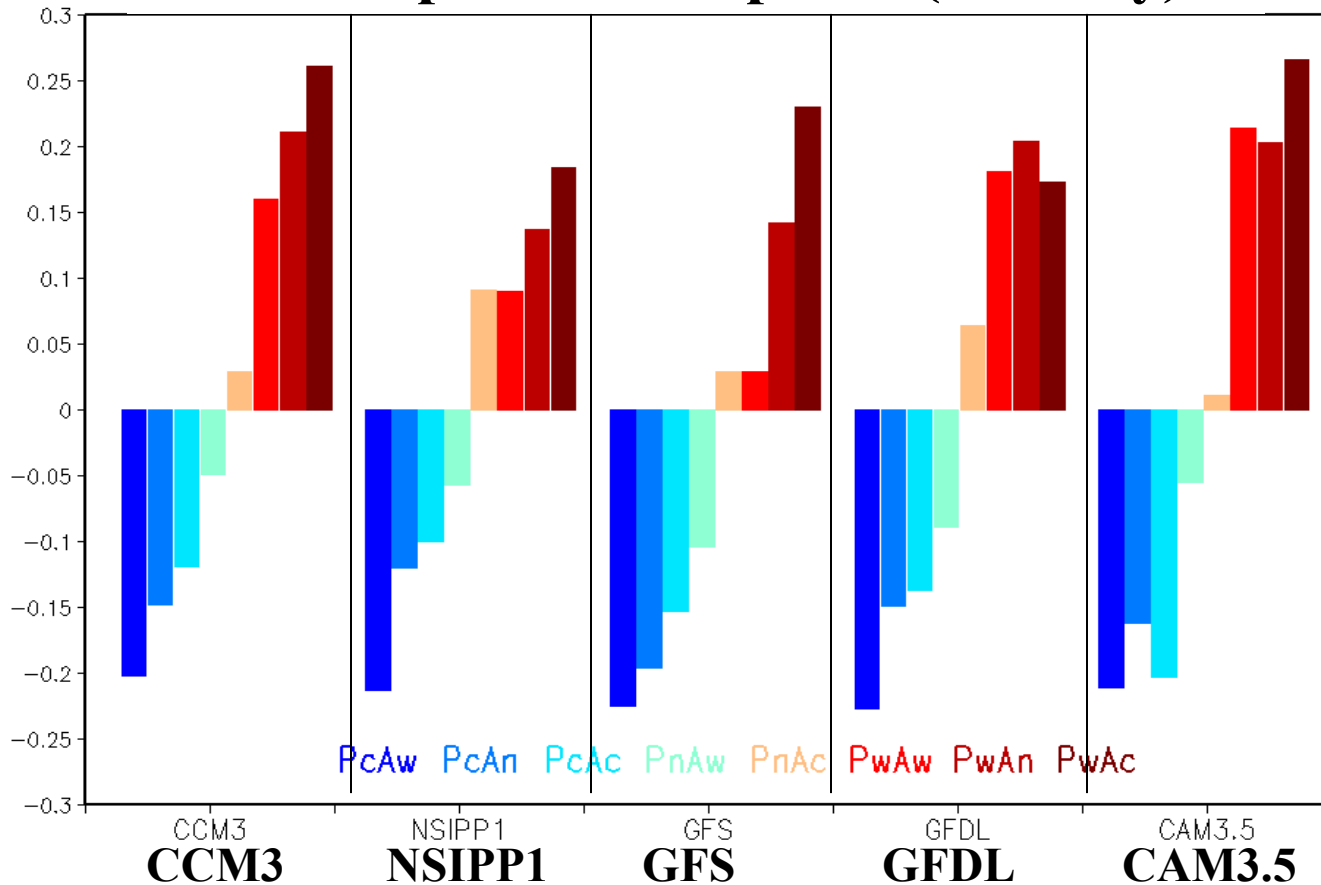


Atlantic
Pattern
(Atl)



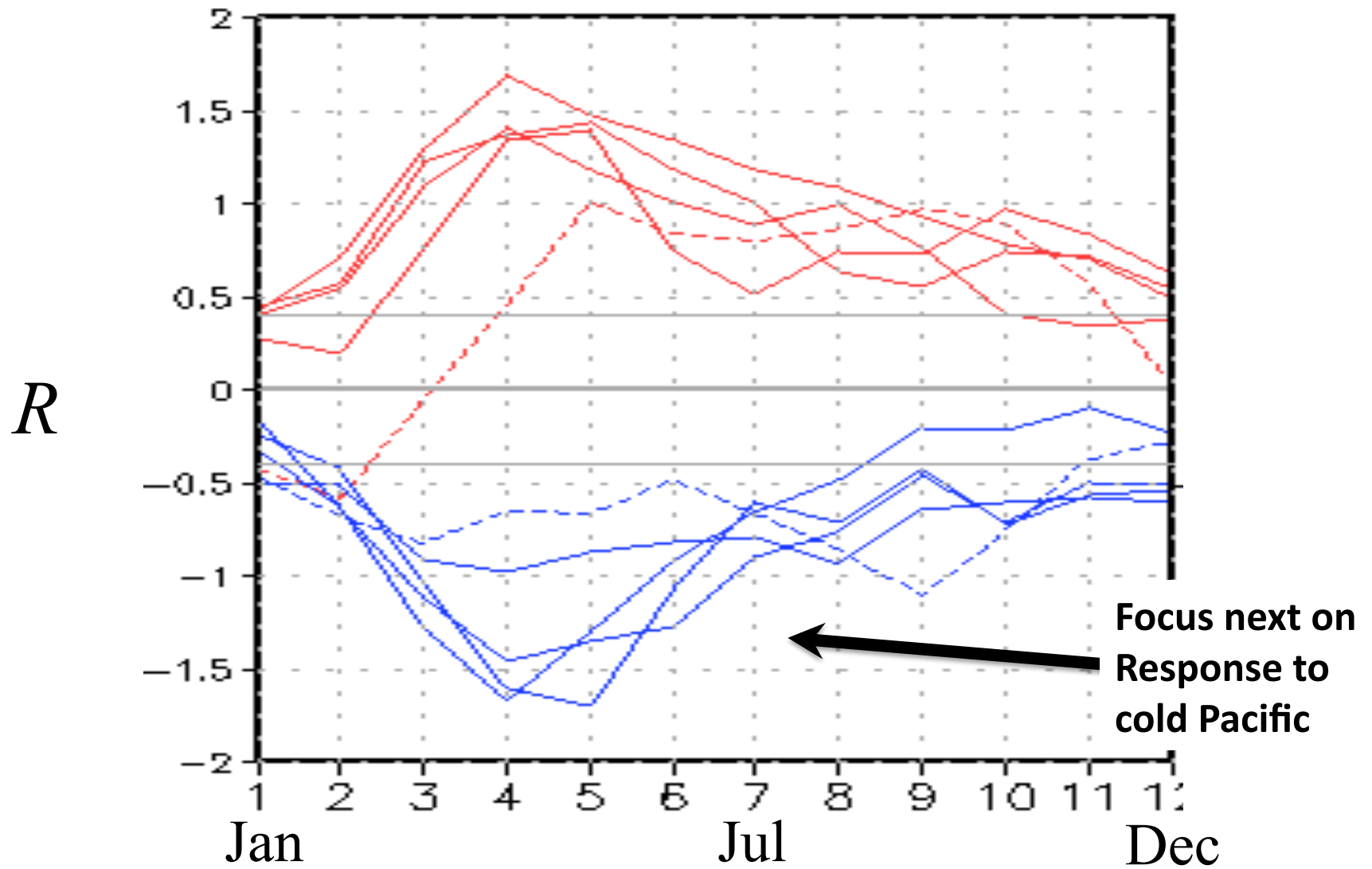
$$(\bar{x} - \bar{y})$$

U.S. Precipitation Response (mm/day)



The annual and continental United States mean responses for precipitation for all 8 combinations of the Pacific and Atlantic SST patterns for the 5 AGCMs. PcAw

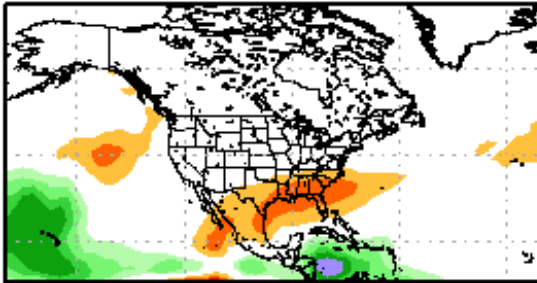
Seasonality of Sig-to-Noise Ratios of US Precipitation for each of the 5 AGCMs in Response to to the **Warm** and **Cold** Pacific SST Patterns



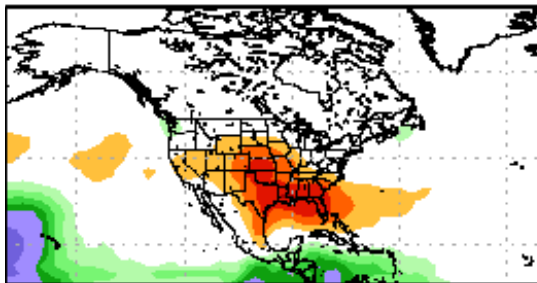
Response to Cold Pacific: Precipitation

Model Response (Signal)

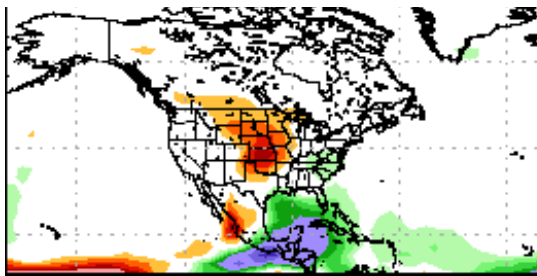
DJF



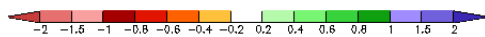
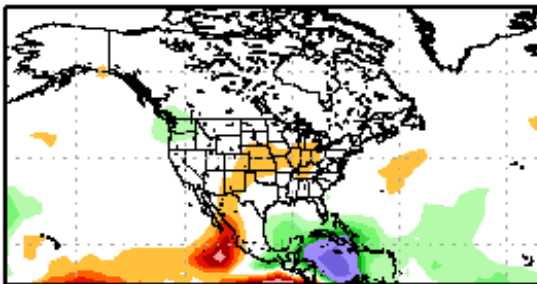
MAM



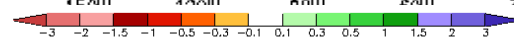
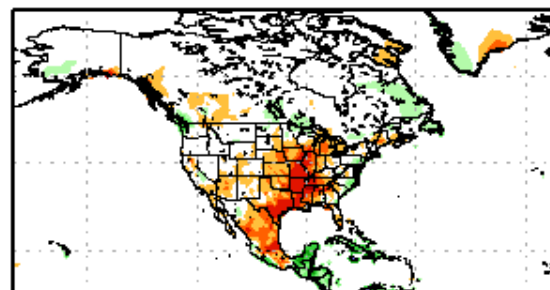
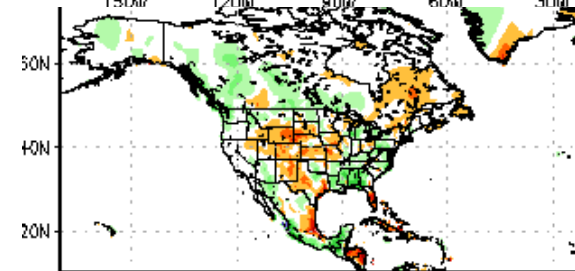
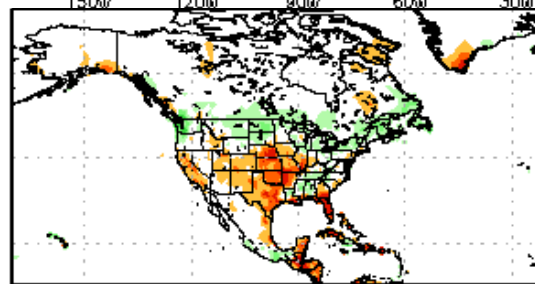
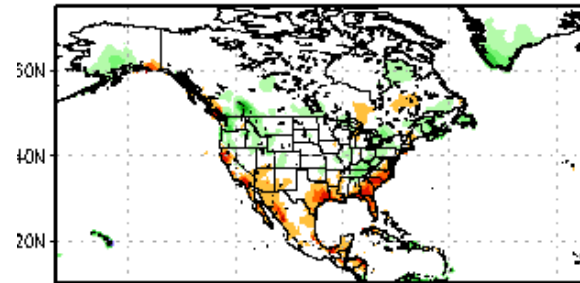
JJA



SON



Cold Pacific Composites (< -1 STD): HadCRU Precipitation Data (1901-2004)



mm/day

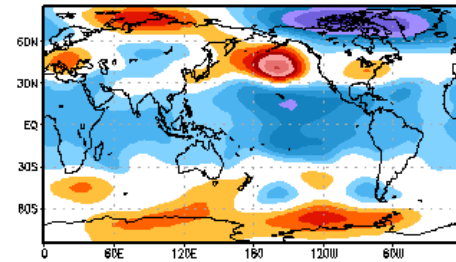
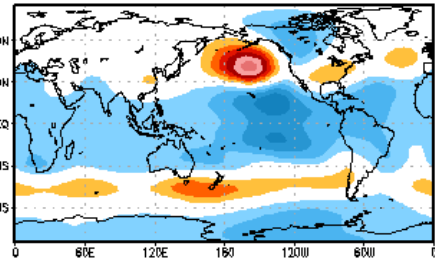
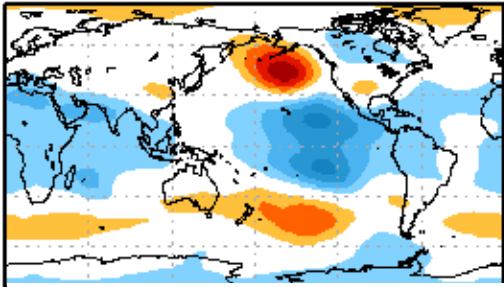
DJF to MAM Transition: Response to Cold Pacific: 200mb Height

Robust Multi-Model Response
To Idealized SST

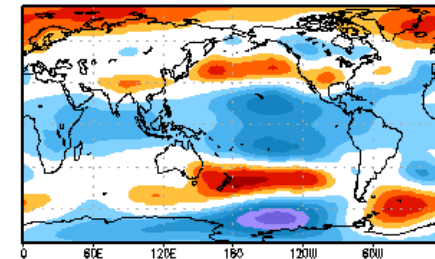
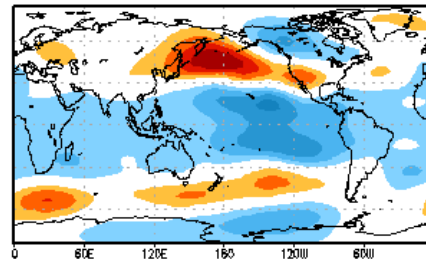
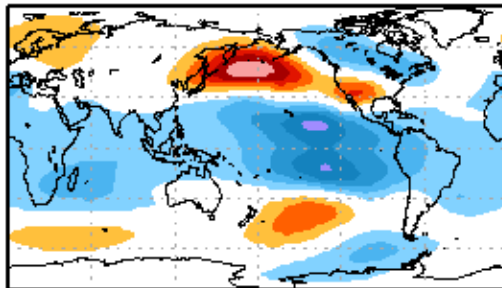
AMIP-NSIPP: Cold
Composite (1901-2004)

Reanalysis-GMAO: Cold
Composite (1948-2004)

DJF

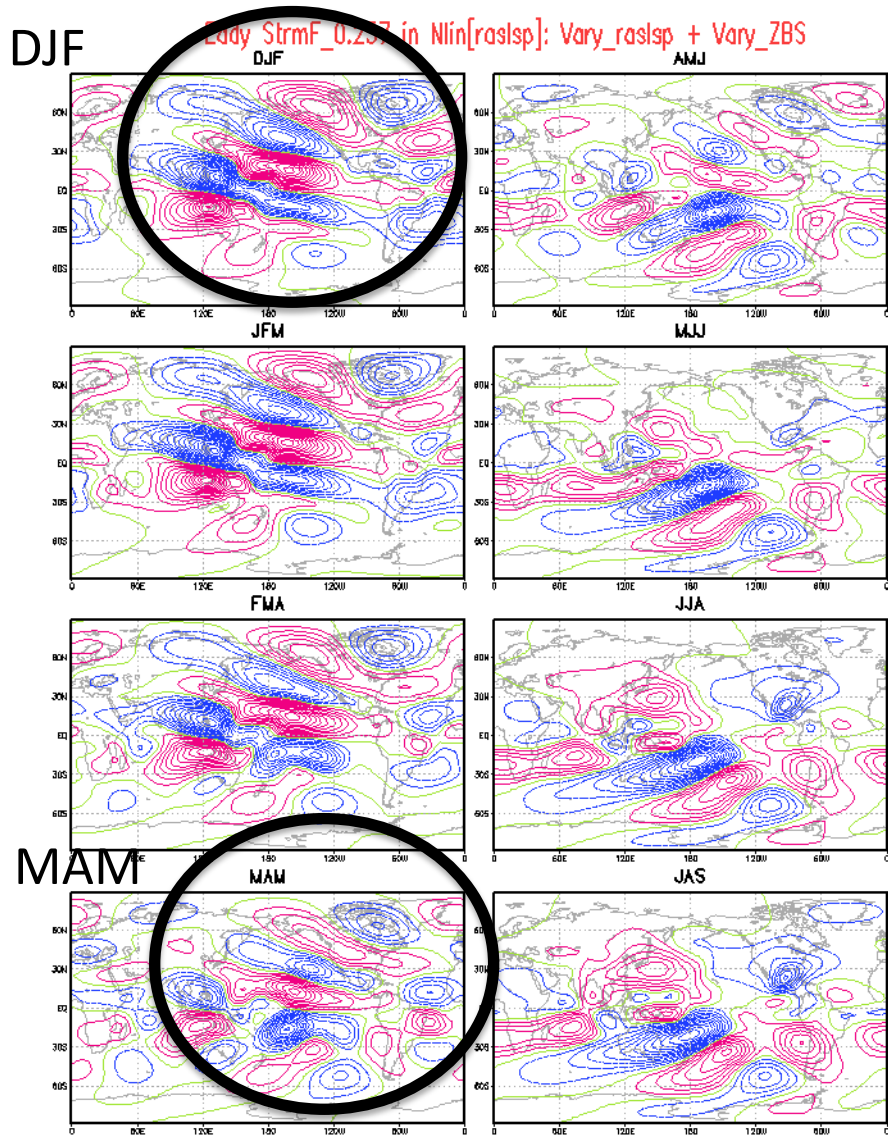


MAM

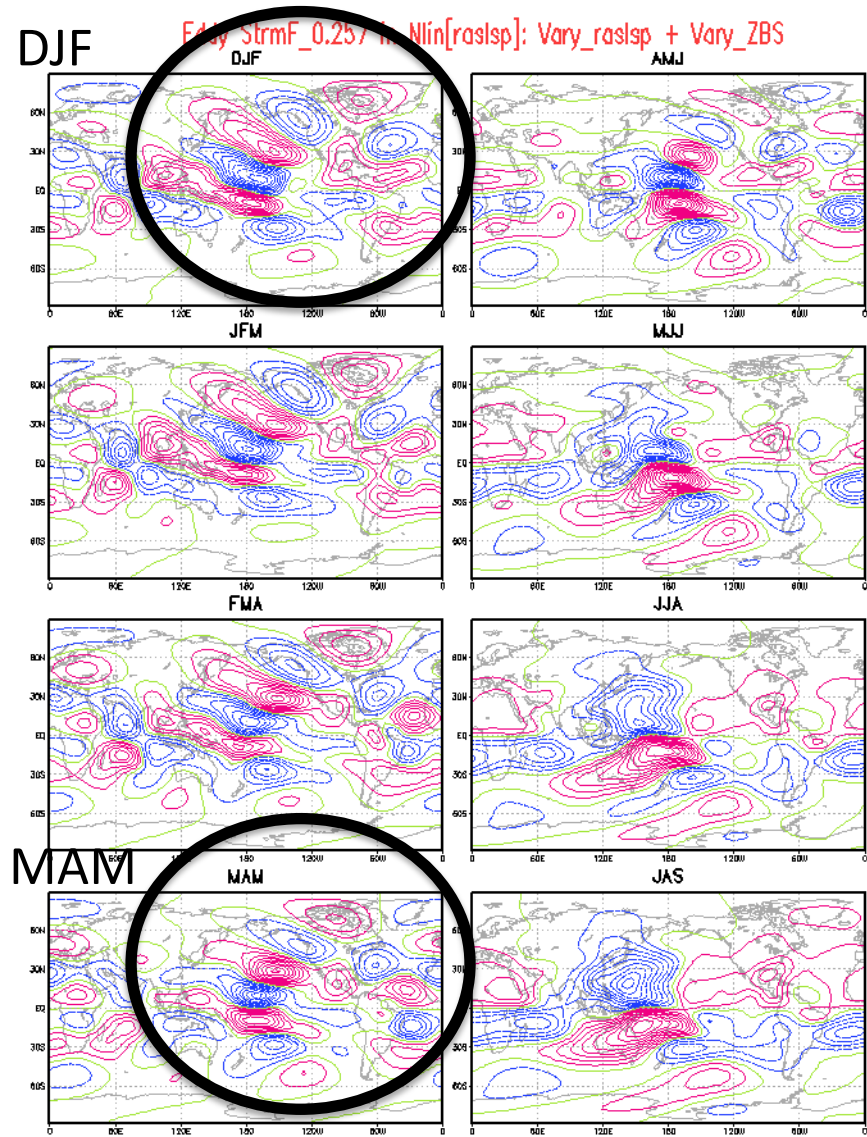


m

Stationary Wave Model Response in Upper Level Stream Function (Zonally-Symmetric Basic State)



Warm Pacific Forcing (from NSIPP Model)



Cold Pacific Forcing (from NSIPP Model)

Key Points- Decadal versus ENSO SST

To the extent that decadal SST persist throughout the annual cycle the transition seasons play a more important role

- Wide spread drought in MAM: a result of the annual cycle of the base state – not a seasonal change in the SST forcing

- Fall also plays an important role: note clear from models why, but likely tied to Walker Circulation response

Separating decadal drought from trends

*Attribution of the seasonality and regionality in climate trends over the United States during 1950-2000". Wang, H., S.D. Schubert, M. J. Suarez, J. Chen, M. Hoerling, A. Kumar and P. Pegion, J. Climate, 2571-2590, 2009.

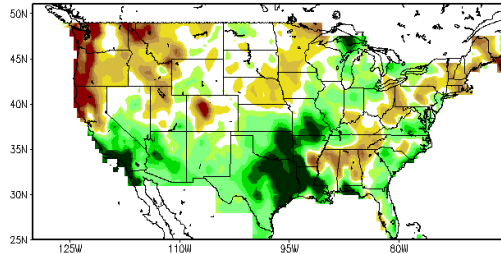
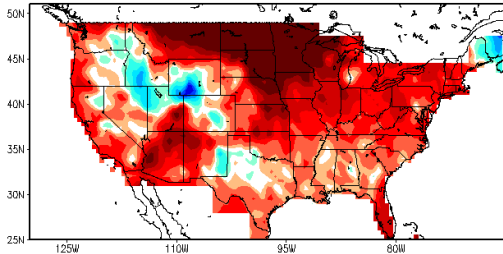
Observed Seasonality and Regionality in Climate Trends over U.S.

HadCRU; 1950-2000

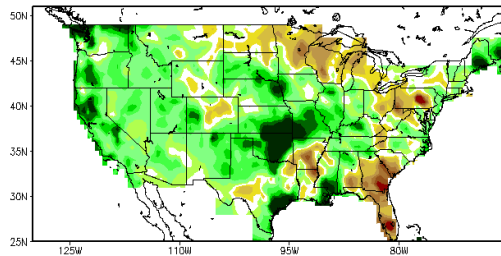
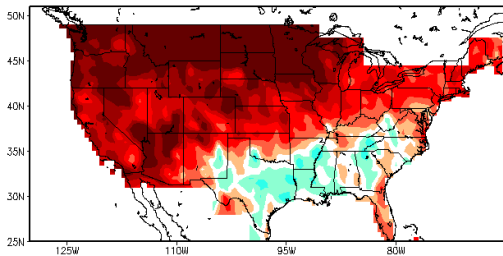
Surface Air Temperature

DJF

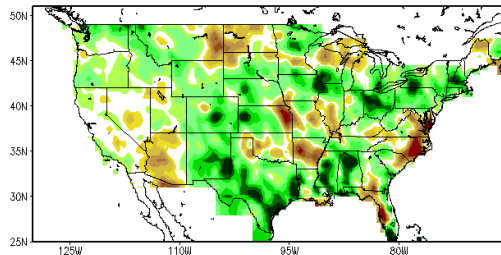
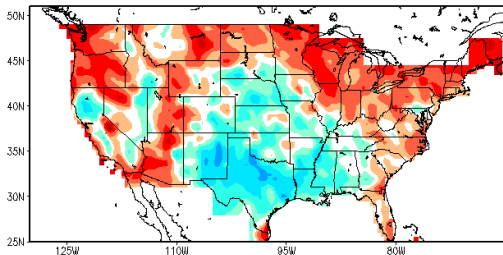
Precip



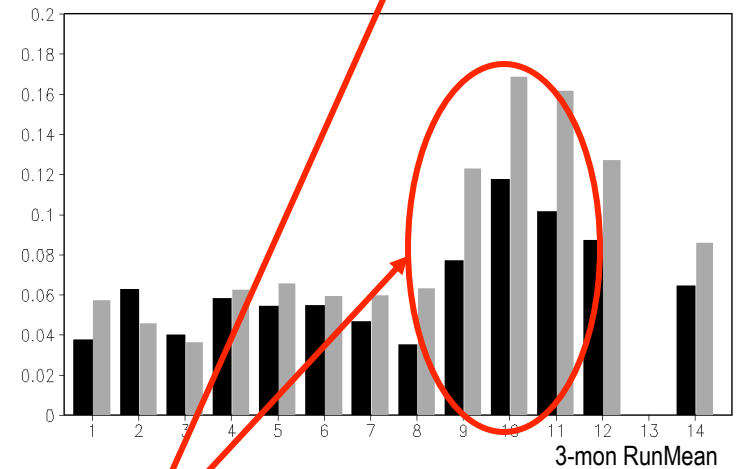
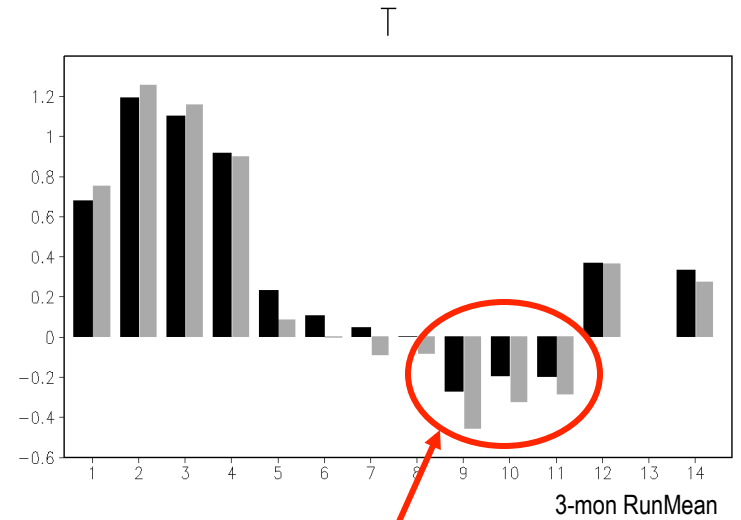
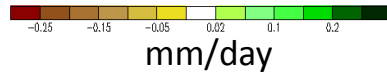
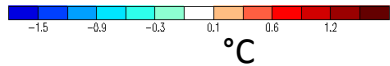
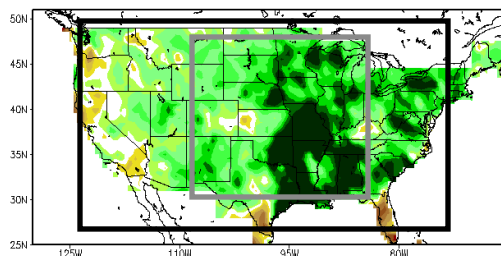
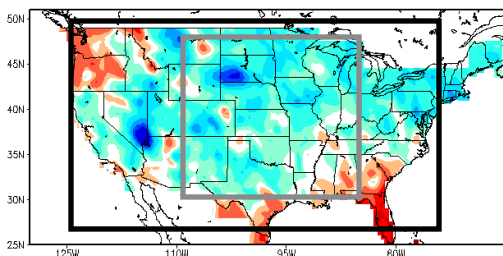
MAM



JJA



SON



Black: US mean (235°E-285°E; 26°N-50°N)
 Gray: central US mean (250°E-275°E; 30°N-48°N)

Distinct cooling and wetting trends during late summer and fall

Model (AMIP) vs Obs

NASA NSIPP AMIP EnsMean(14)

HadCRU

Surface Air Temperature

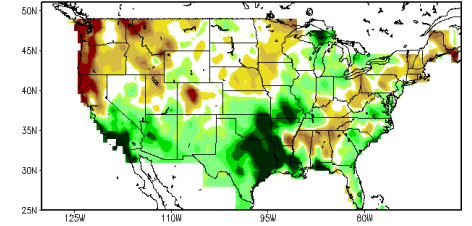
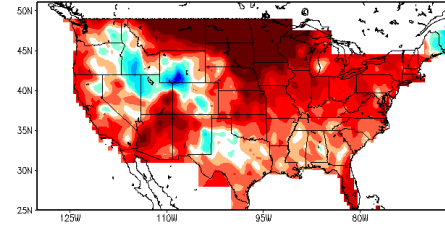
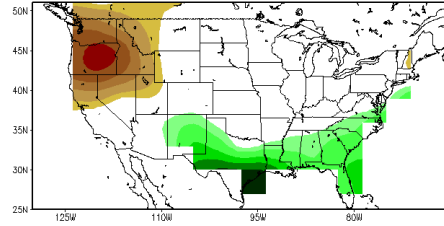
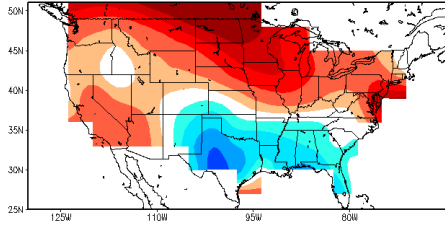
DJF

Precip

Surface Air Temperature

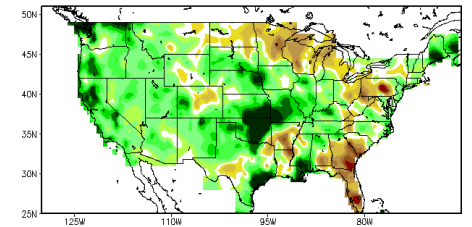
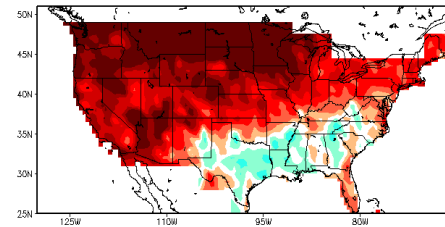
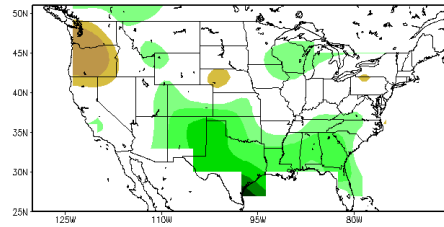
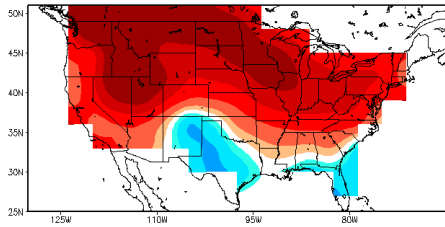
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Precip



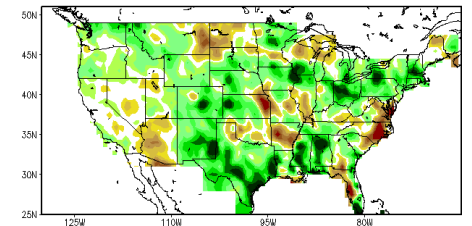
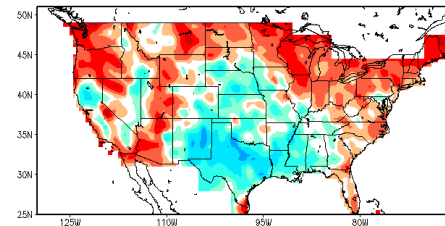
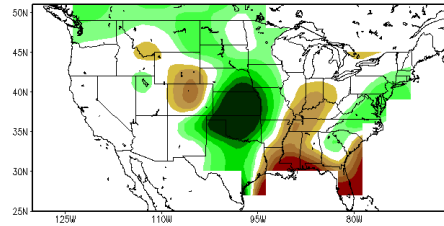
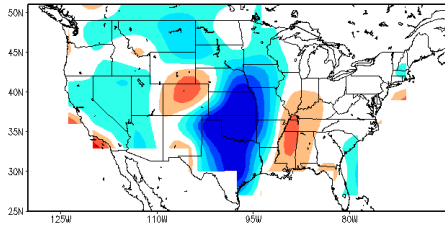
MAM

MAM



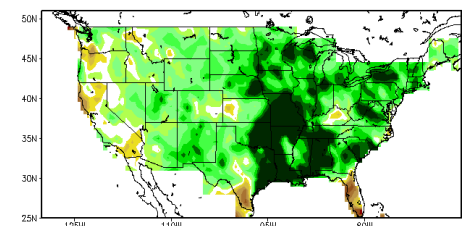
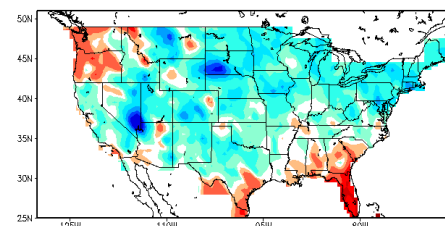
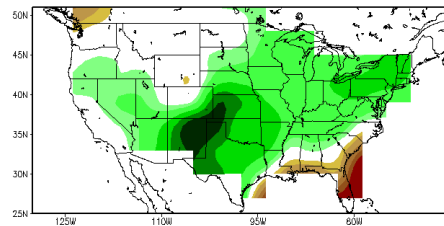
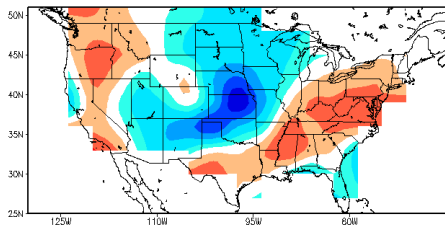
JJA

JJA

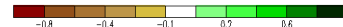


SON

SON



°C



mm/day



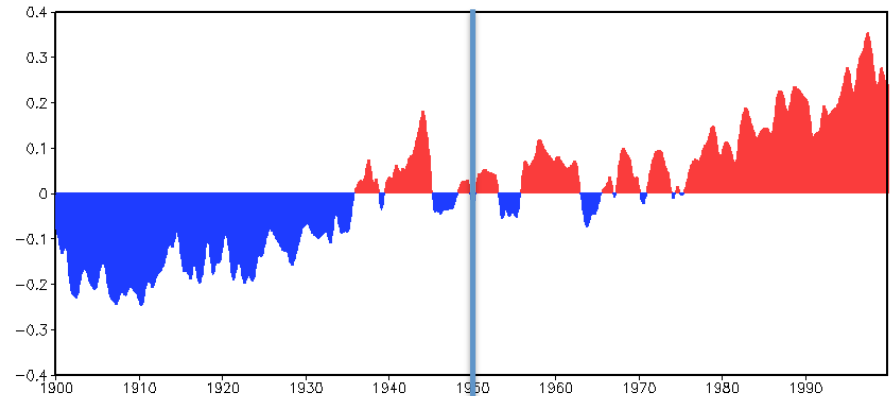
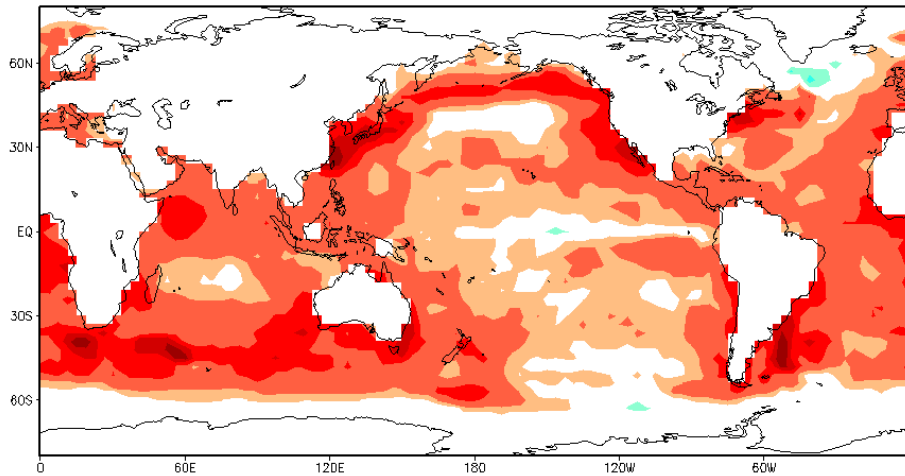
°C



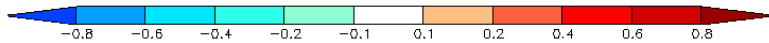
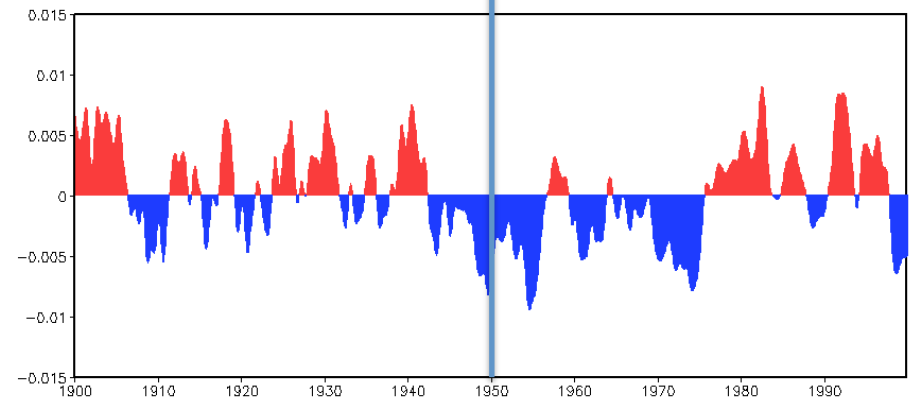
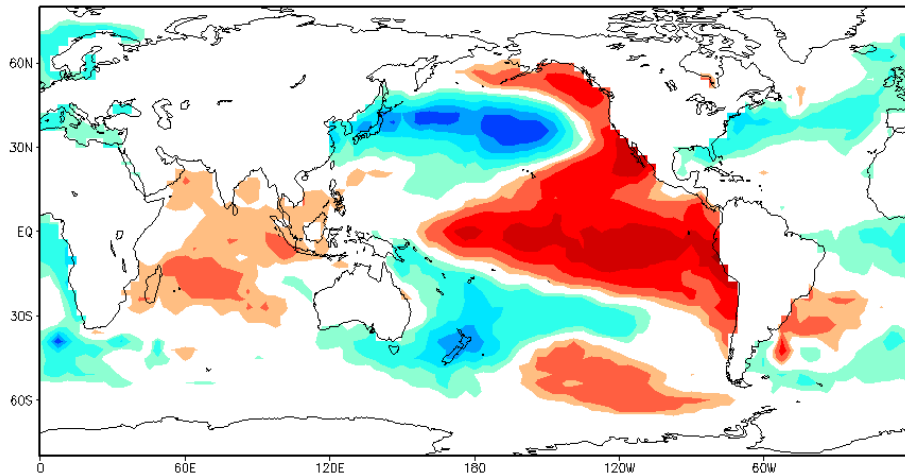
mm/day

Leading SST EOFs: Global Warming (GW) and Pacific Decadal Variability (PDV) Hadley SST; 1901-2000

Global Warming Mode (GW) 32%



Decadal Variability Mode (DV) 12%

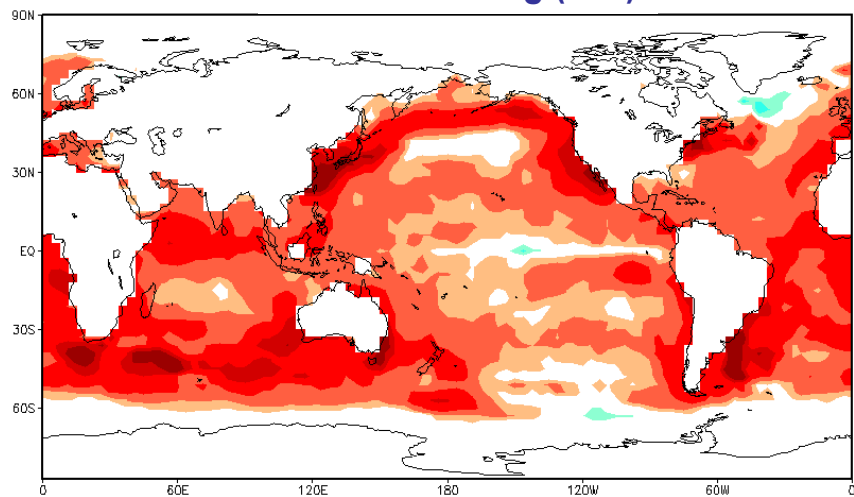


Chen *et al.* (2007)

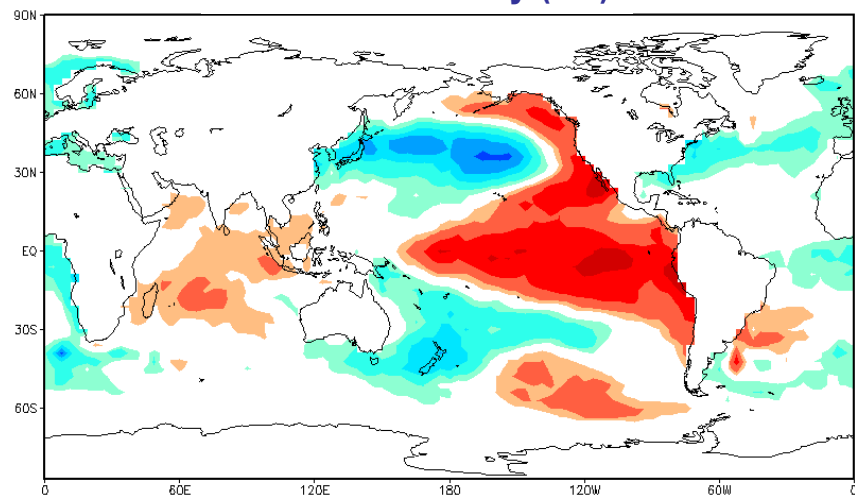
Linear trend of SST: Total=GW+DV+Residual

Hadley SST; 1950-2000

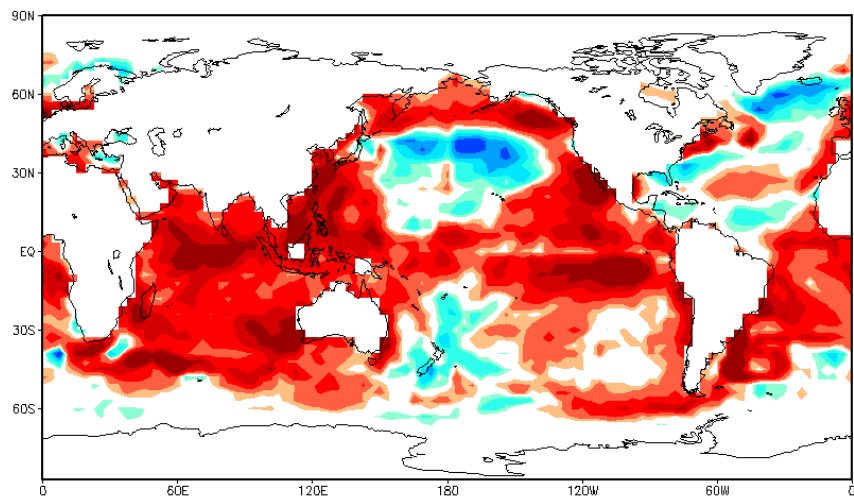
Global Warming (GW)



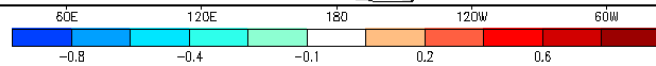
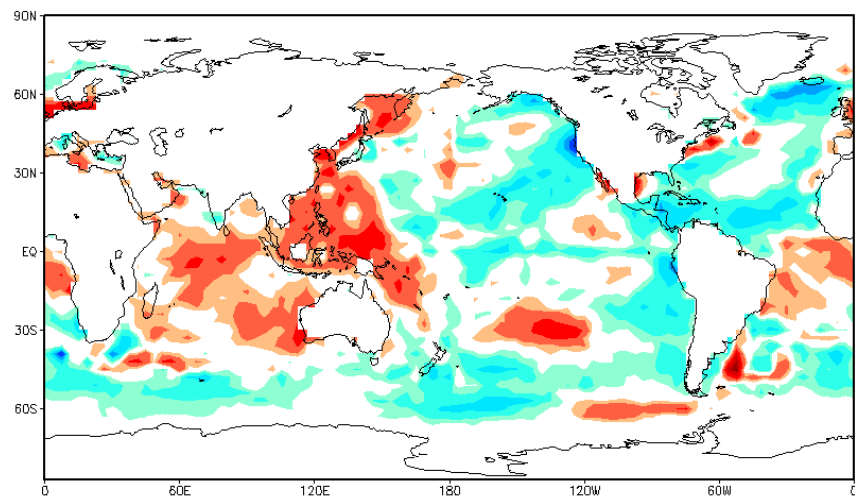
Decadal Variability (DV)



Total



Residual



Input for a set of idealized AGCM experiments

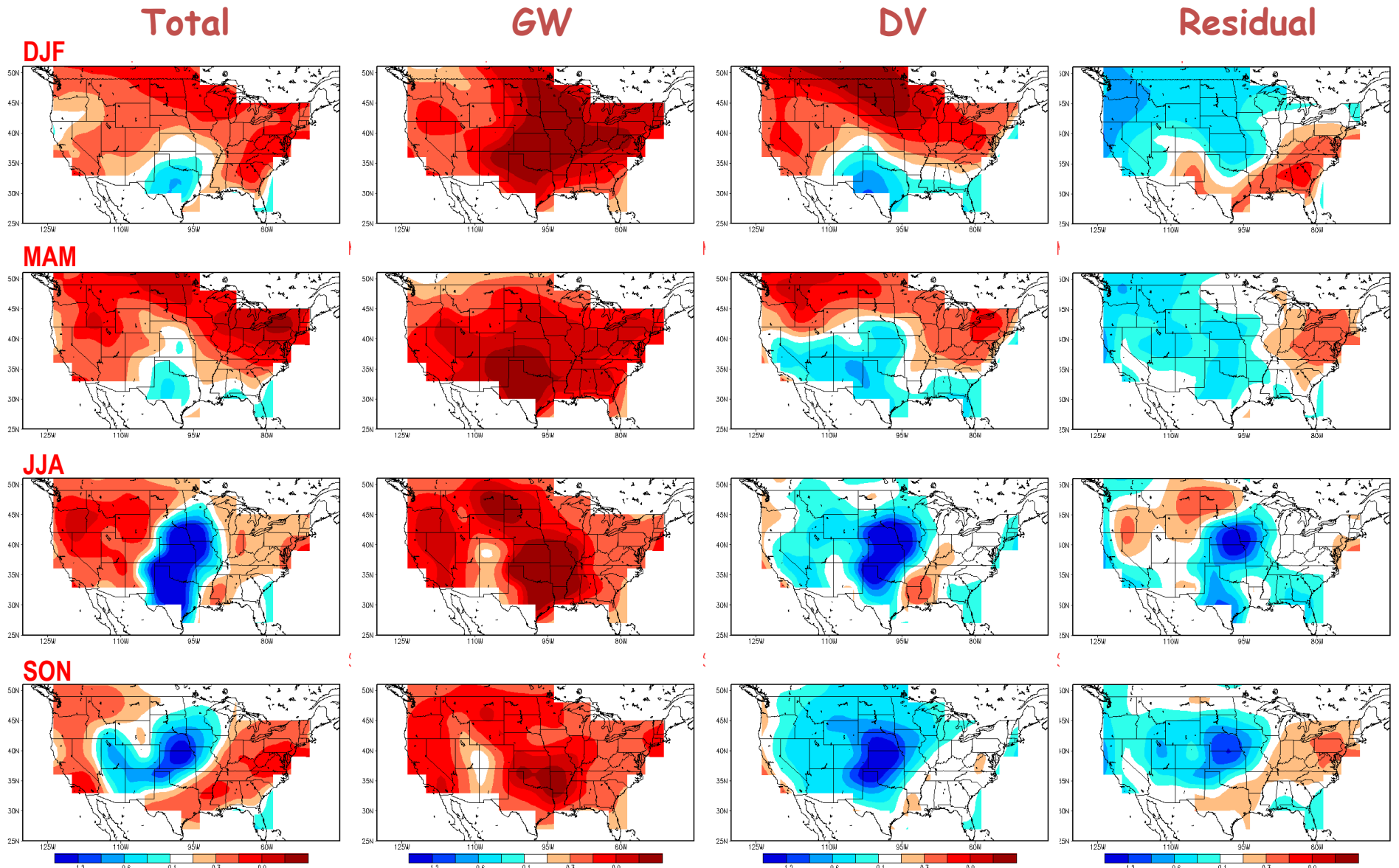
Idealized AGCM Experiments

NASA NSIPP-1 AGCM; 3deg

Experiments	SST
Control	SSTClim(monthly SST climatology over 1944-1976)
Total	SSTClim+SSTA_Total(linear change of SST over Jan1950-Dec2000)
Global Warming (GW)	SSTClim+SSTA_GW(linear change of SST associated with GW EOF)
Uniform Warming	SSTClim+0.32K
Decadal Variability (DV)	SSTClim+SSTA_DV(linear change of SST associated with DV EOF)
DV in Pacific	SSTClim+SSTA_DV_Pac(linear change of SST associated with DV in Pacific)
Residual	SSTClim+SSTA_Residual(SSTA_Total - SSTA_GW - SSTA_DV)

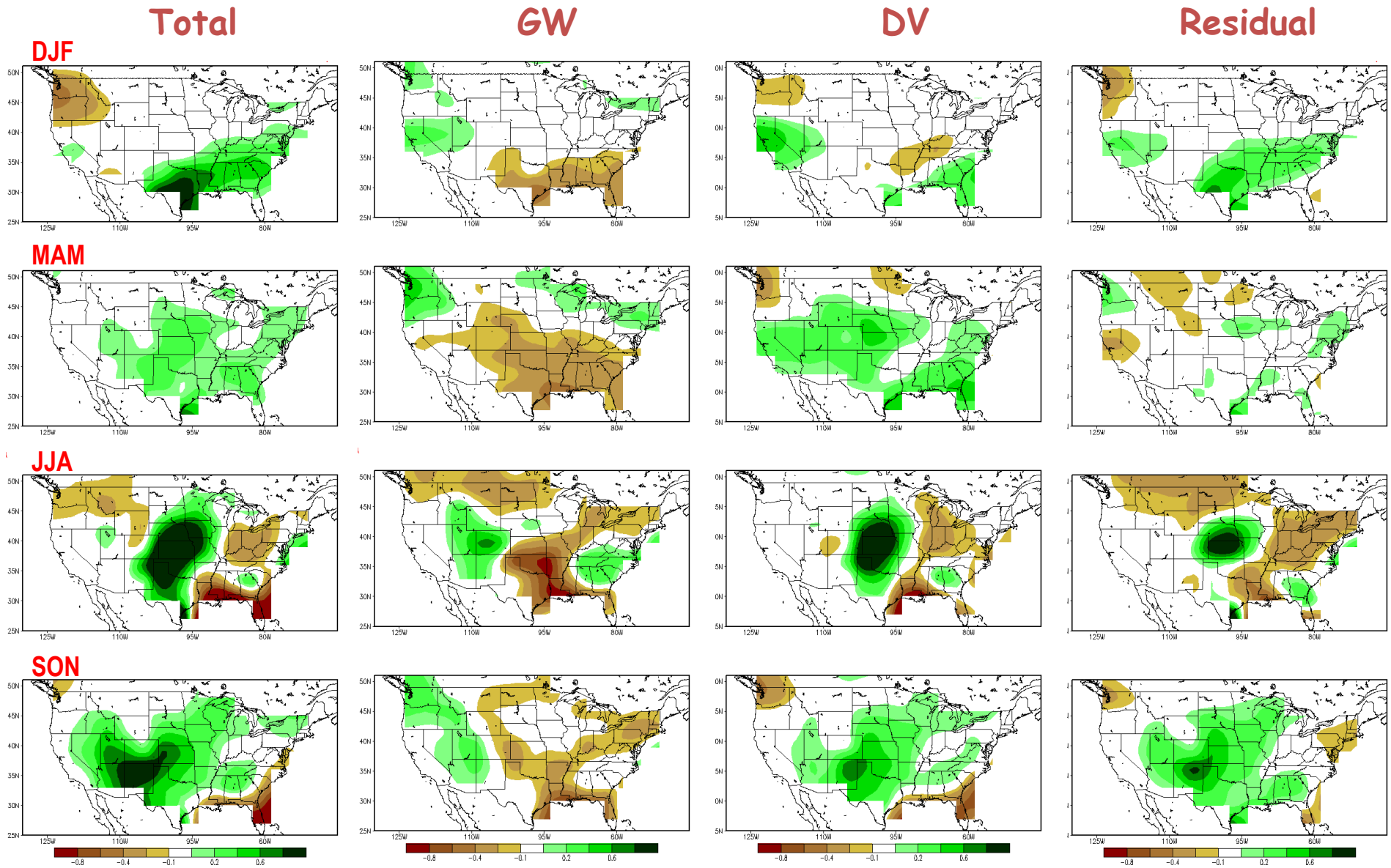
- All runs are integrated for 100 years
- Data averaged over the last 60 years taken as climatology
- Climatological difference between control run and an anomaly run represents effect of corresponding SST trend

Surface Air Temperature (T) Response in Idealized AGCM Experiments



°C

Precip (P) Response in Idealized AGCM Experiments



mm/day

Conclusions

- The observed climate trends over the US during 1950-2000 can be mostly explained by changes in SST.
- Among the leading SST patterns:
 - Global Warming (GW):
 - mainly leads to a general uniform warming trend
 - Pacific Decadal Variability (PDV):
 - Main contributor to central US cooling and wetting trends in summer/fall
 - Residual (cool Atlantic):
 - mainly contributes in late summer and fall (warming and wetting)

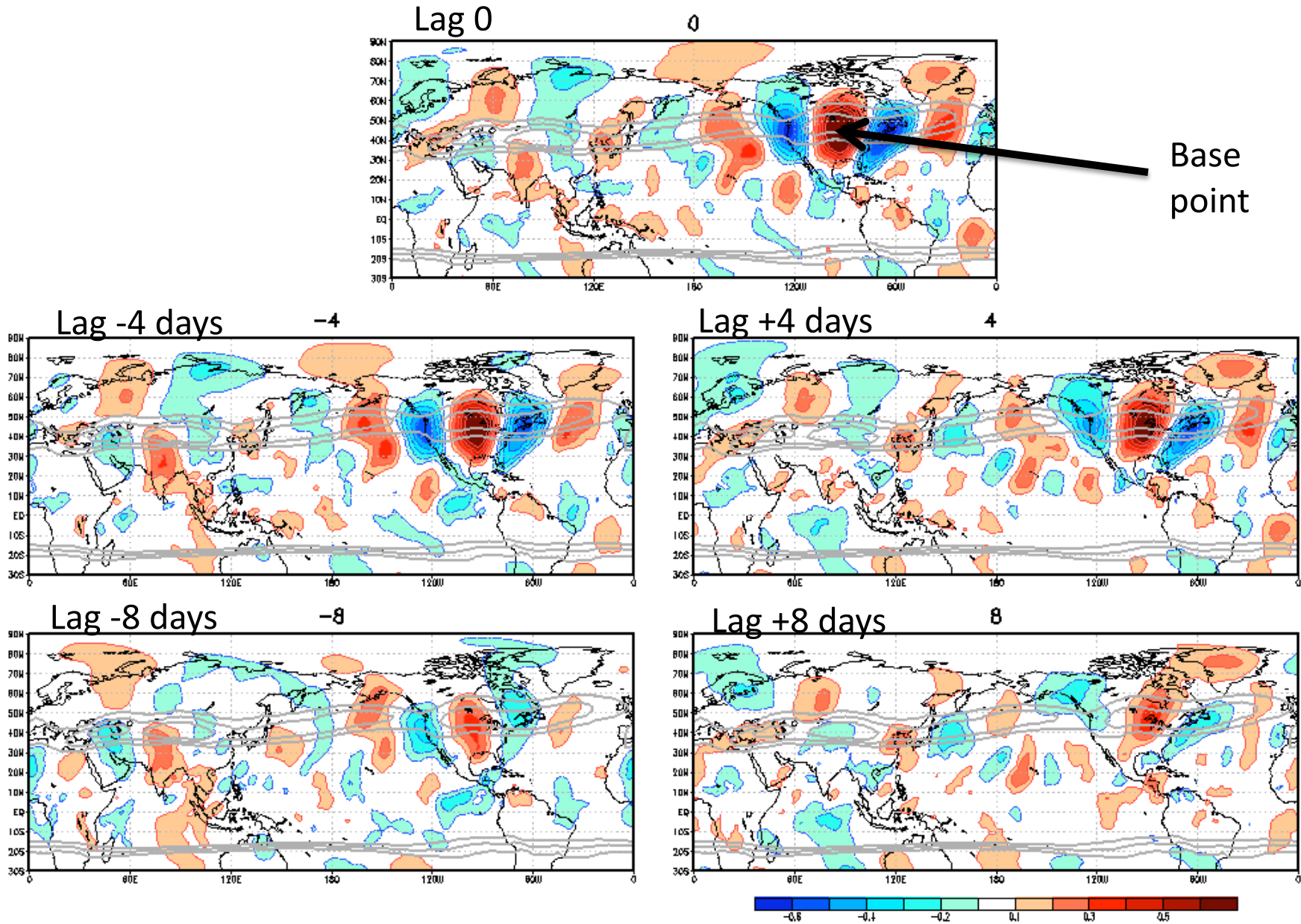
Sub-seasonal drought/heat waves

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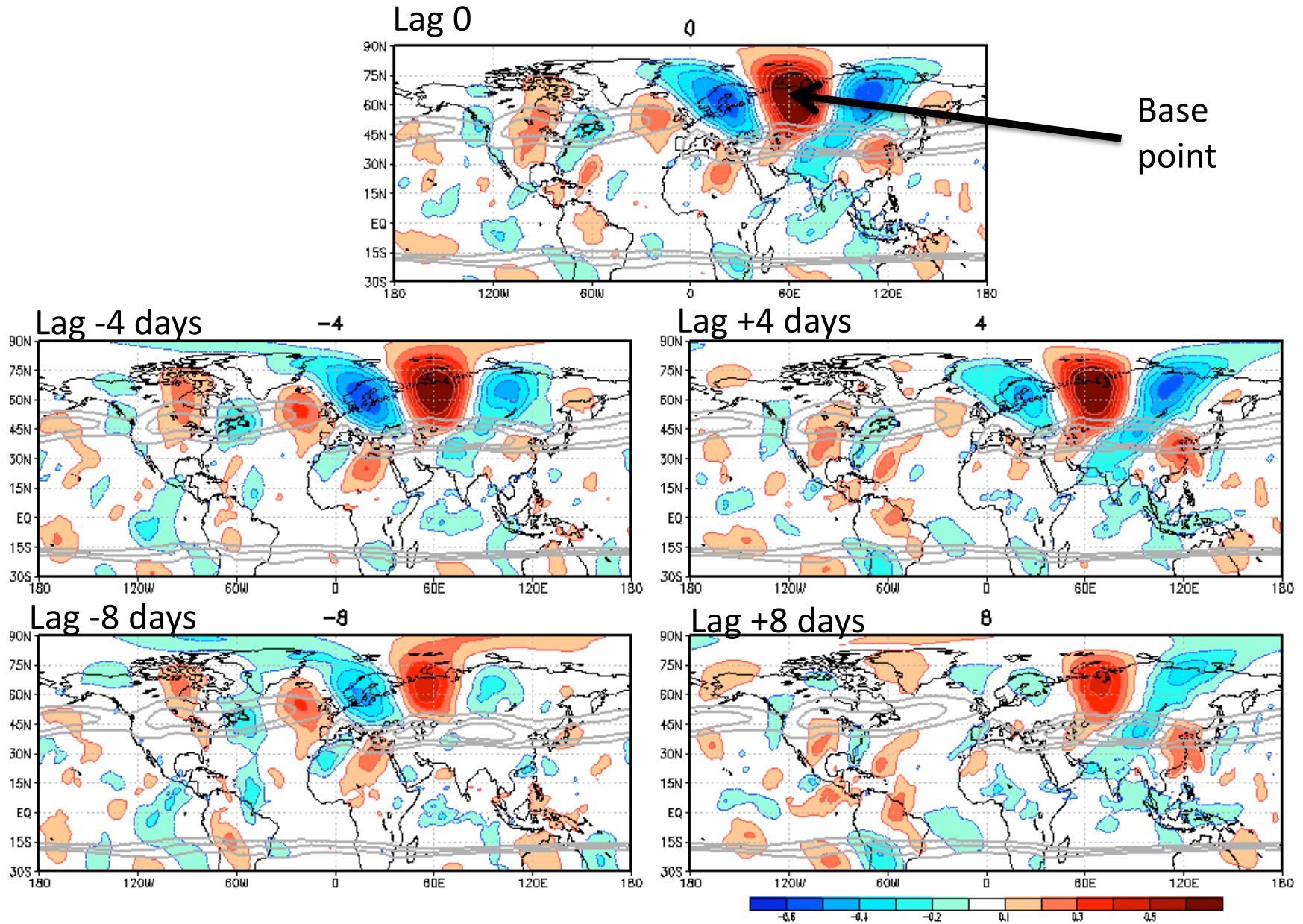
One- point lead/lag Correlation (V250mb)

(30-90 day filter, MERRA - JJA 1979-2008)

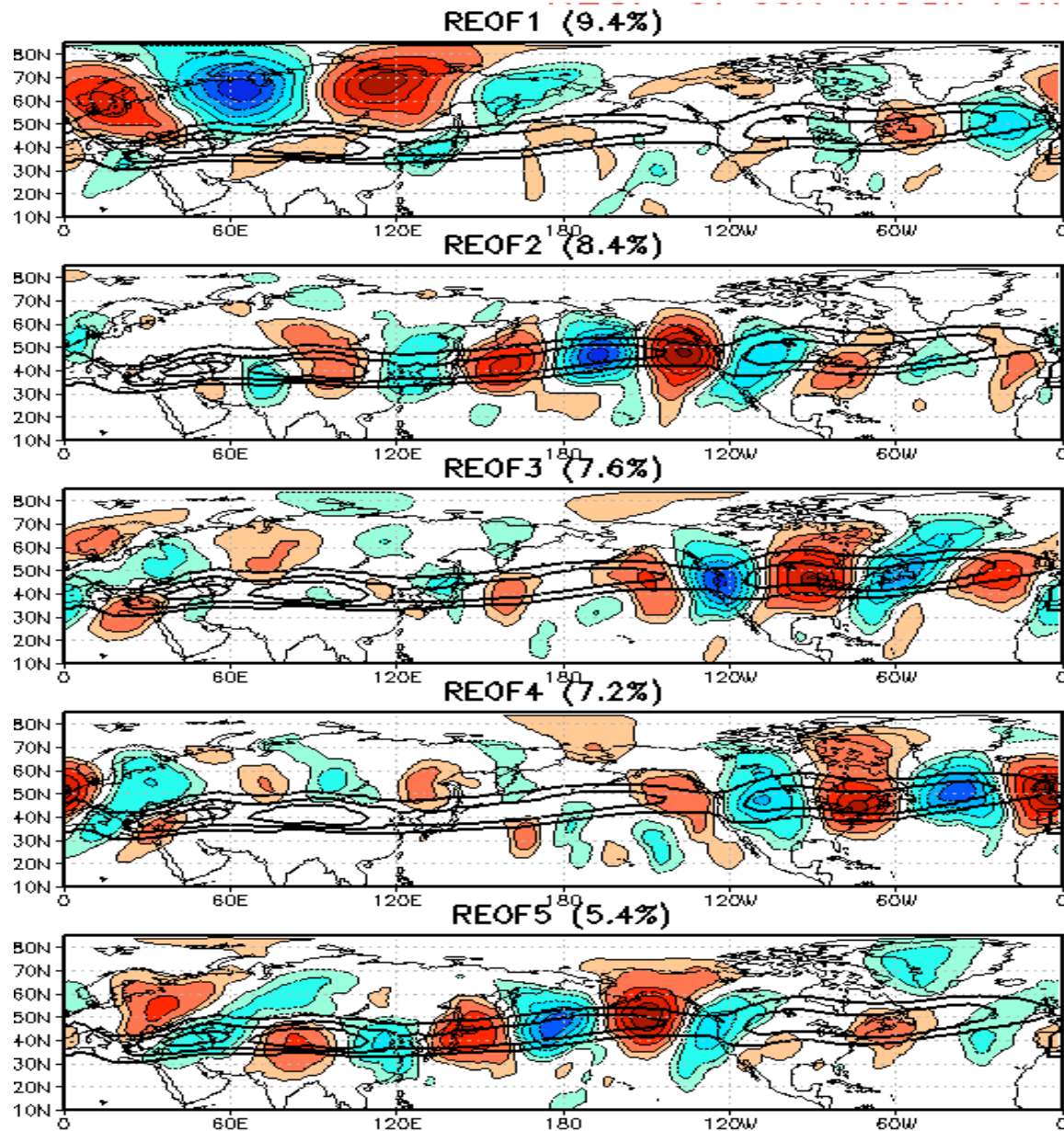


One- point lead/lag Correlation (V250mb)

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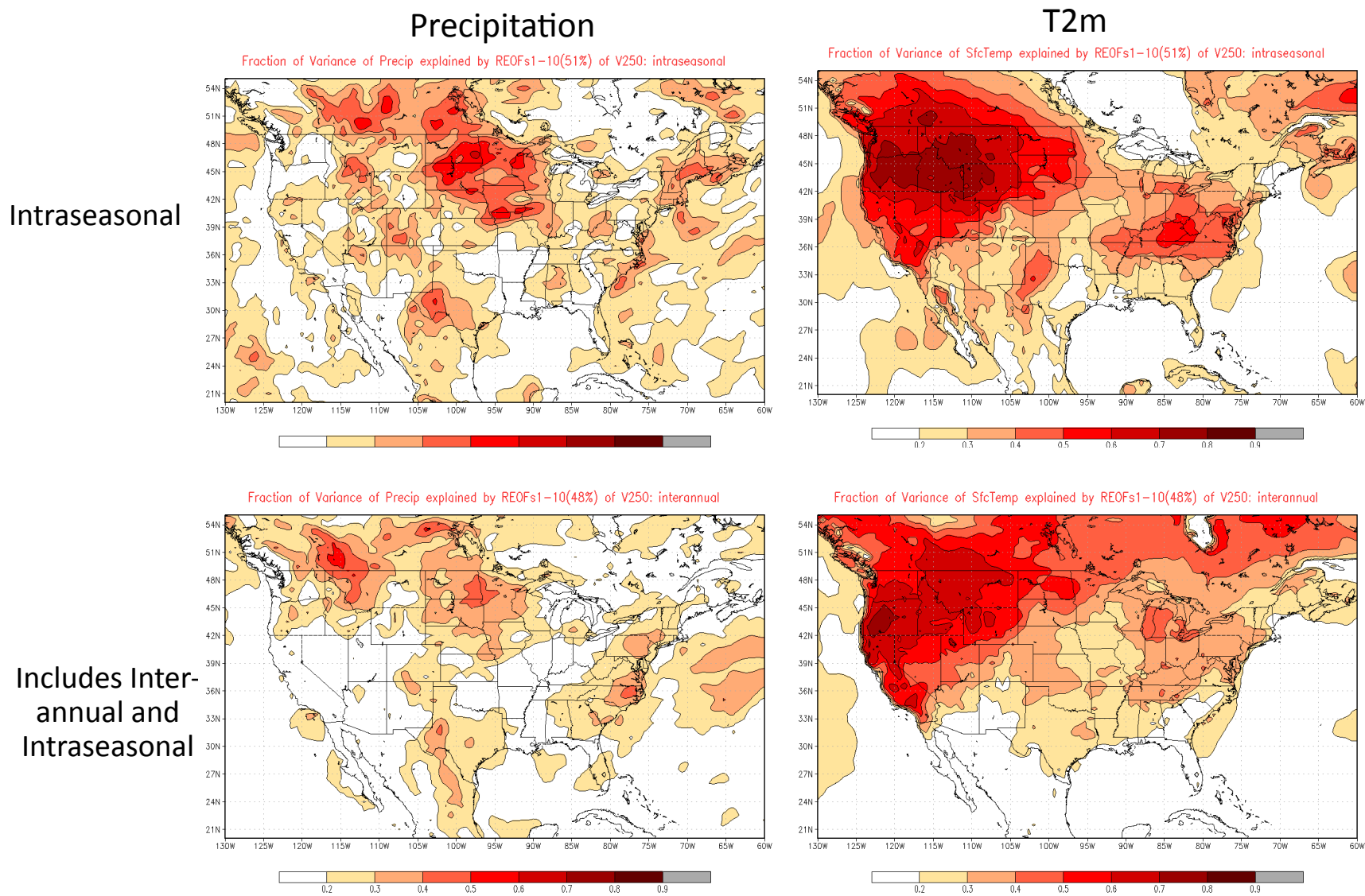


Leading Rotated EOFs of Intraseasonal (Monthly JJA) V250mb



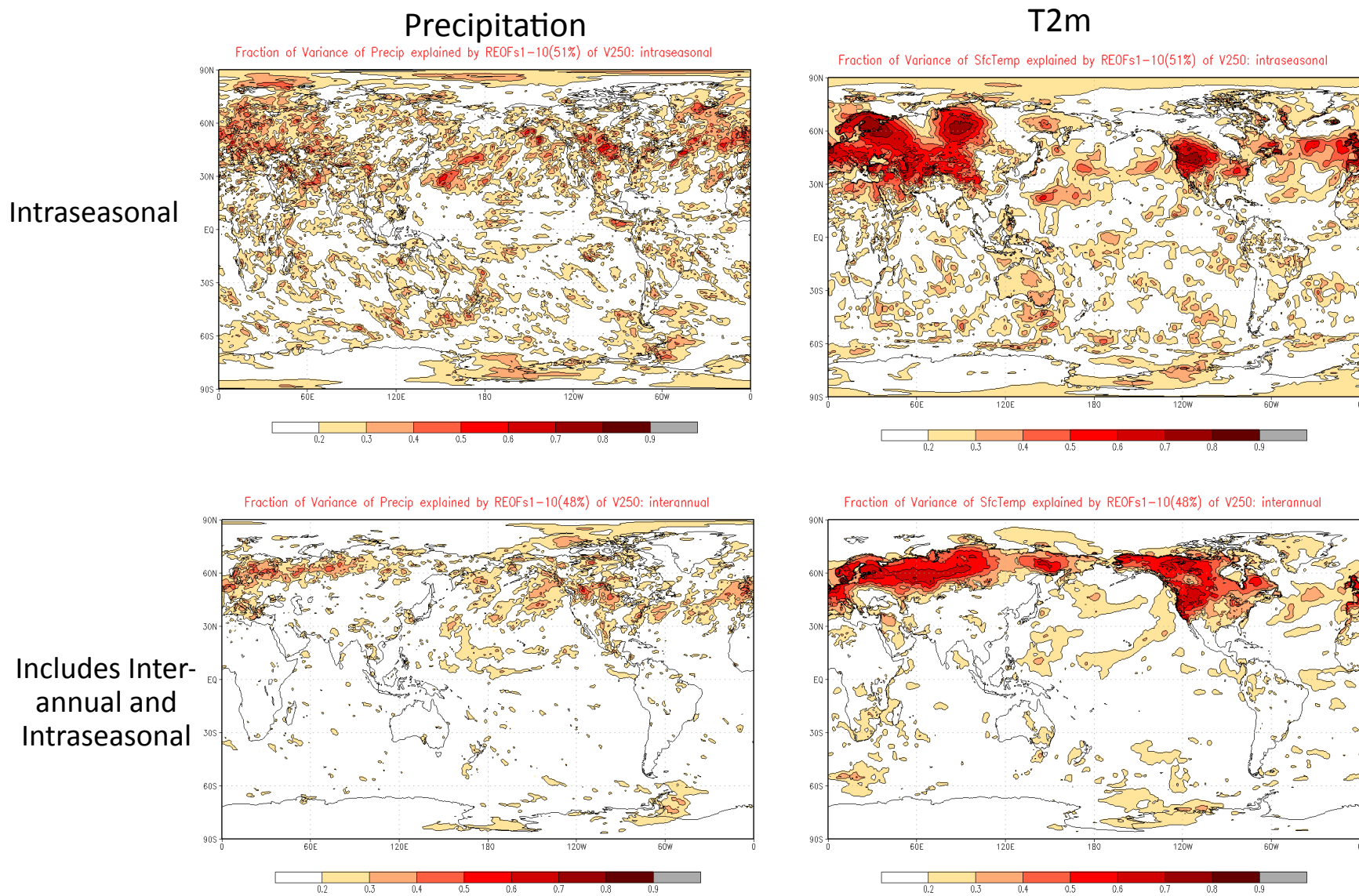
MERRA:
1979-
2008

Fraction of Variance Explained by first 10 REOFs of 250 v-wind



Based on Monthly MERRA data JJA (1979-2008)

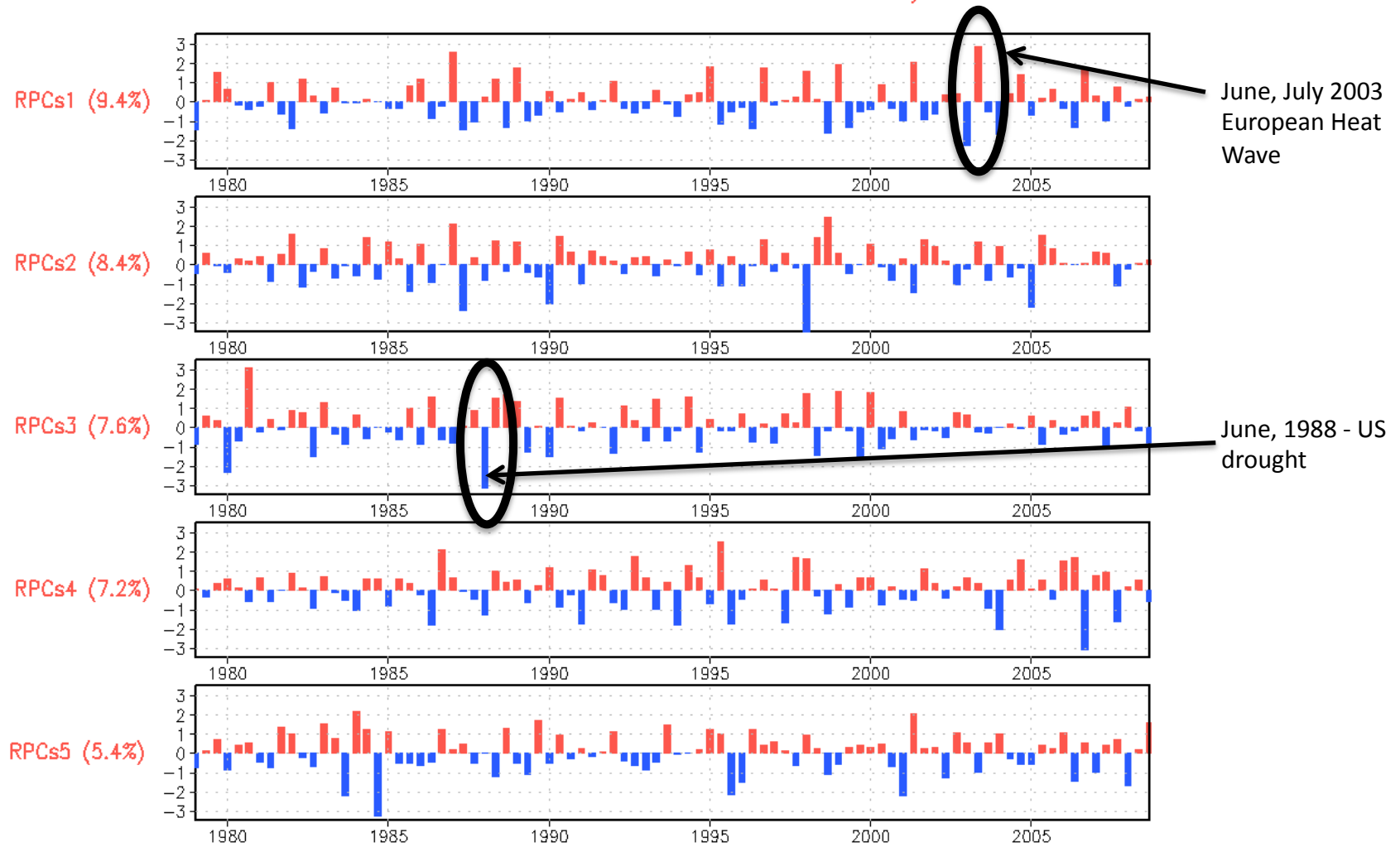
Fraction of Variance Explained by first 10 REOFs of 250 v-wind



Based on Monthly MERRA data JJA (1979-2008)

Leading Normalized Rotated PCs (Intraseasonal Monthly JJA)

Normalized RPCs of JJA mean removed monthly mean v250

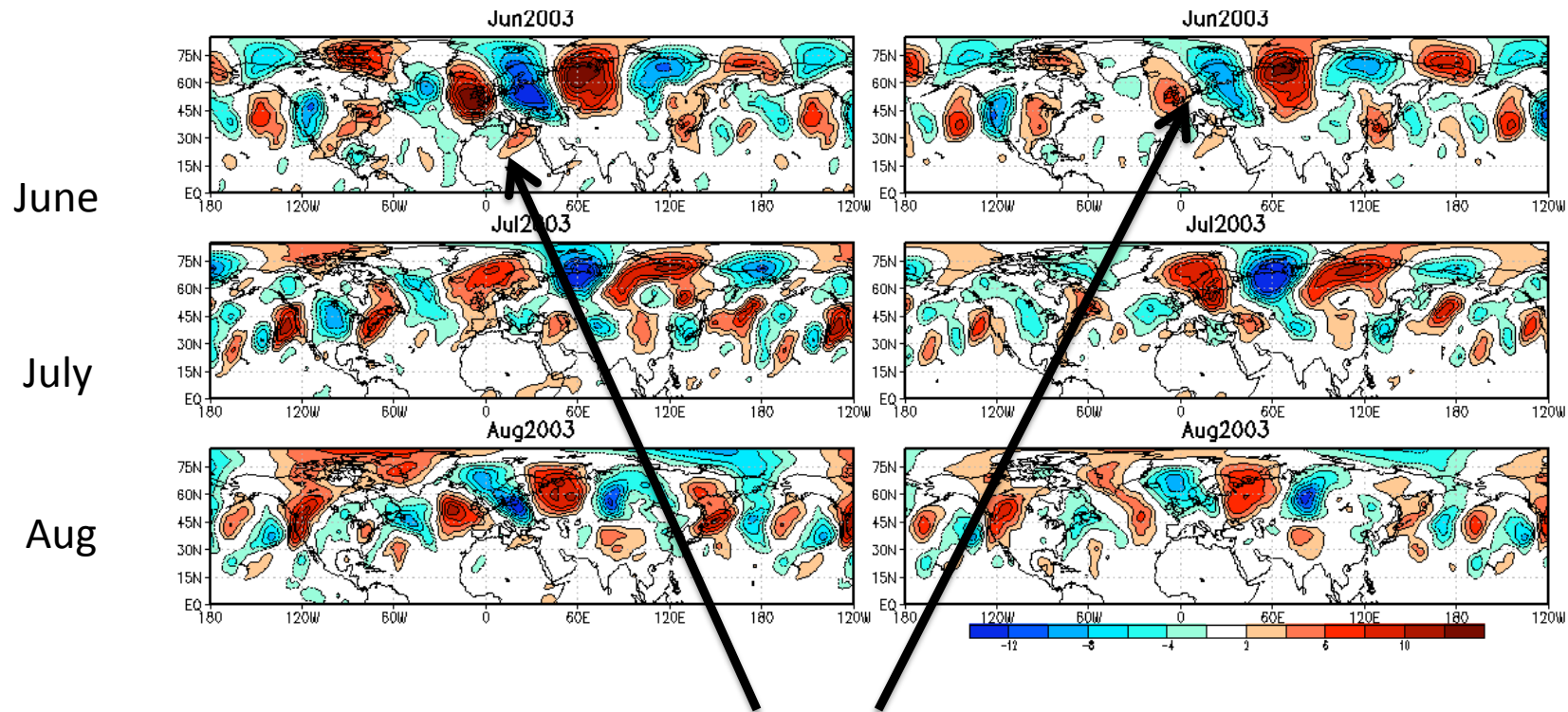


2003 European Heat Wave

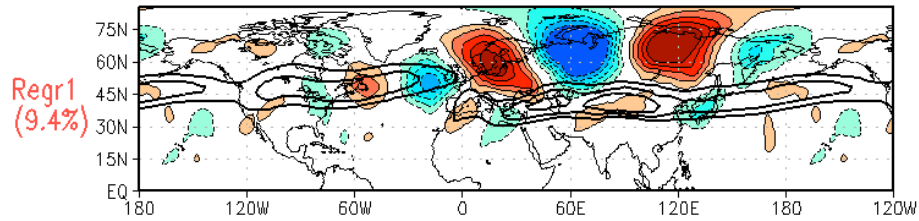
2003 Monthly Anomalies (V250mb)

Interannual

Intraseasonal

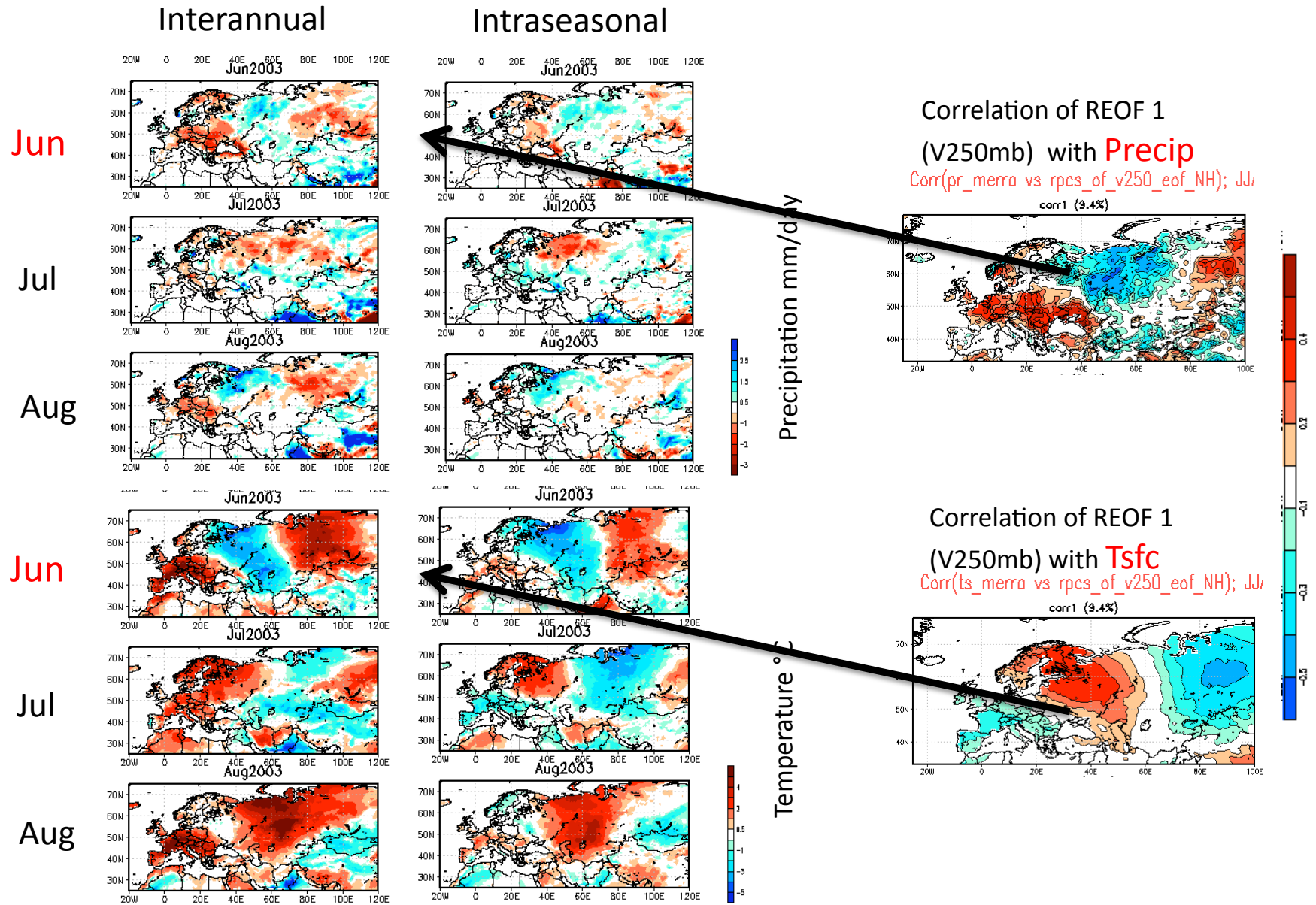


REOF 1 – Regression with V250mb



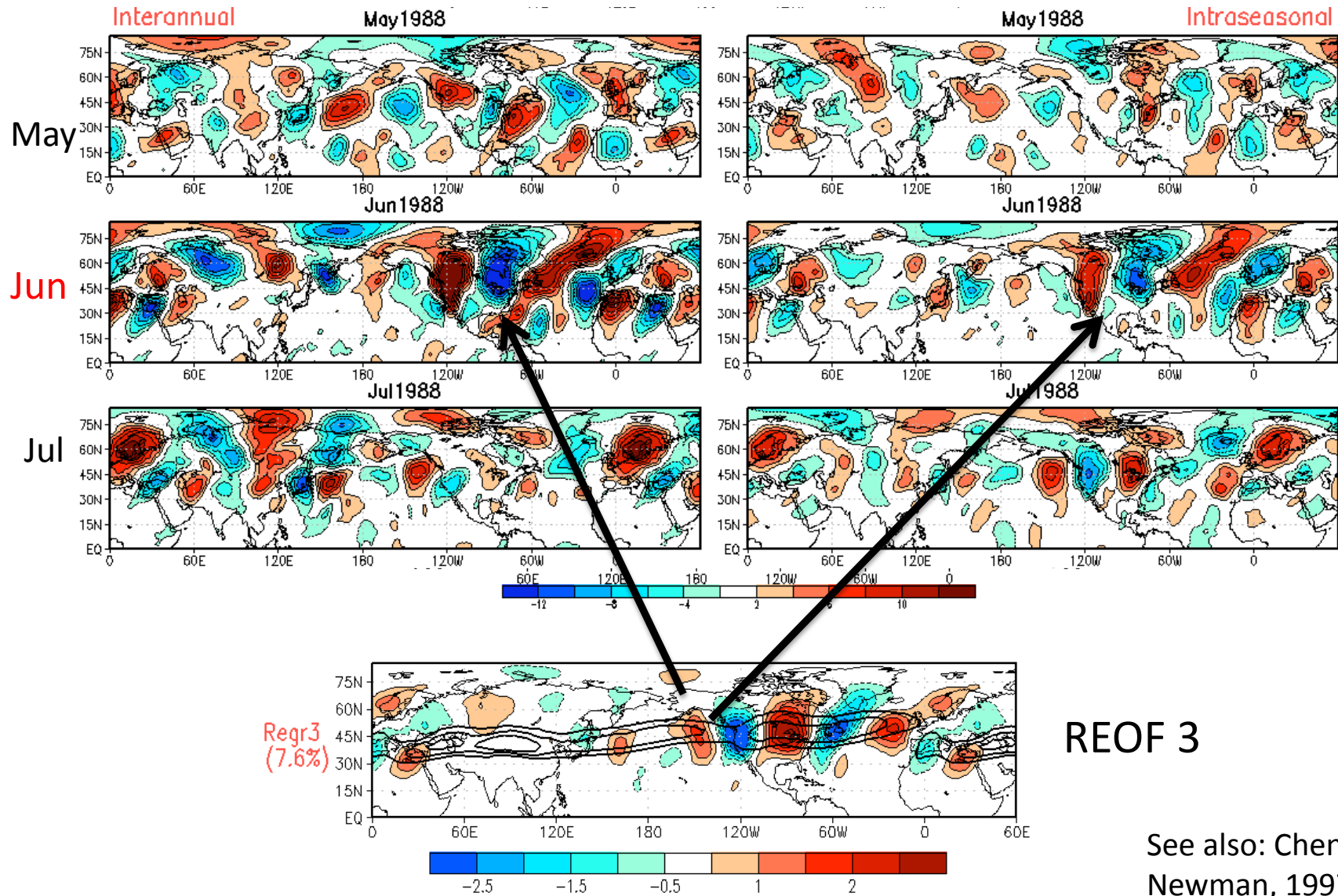
REOF 1

2003 European Heat Wave Monthly Anomalies (MERRA)



1988 US Drought/Heat Wave

1988 Monthly Anomalies (V250mb)



See also: Chen and Newman, 1997

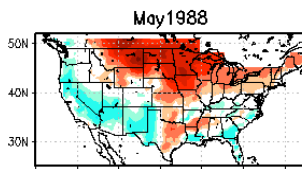
1988 Monthly Anomalies (Tsfc)

Interannual

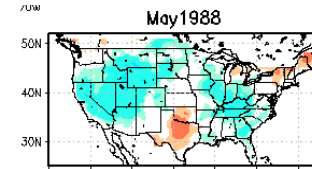
Intraseasonal

Correlation between V250 REOF 3 and Tsfc (MERRA)

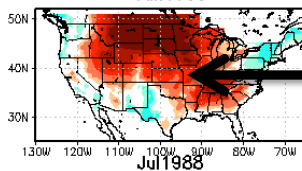
May



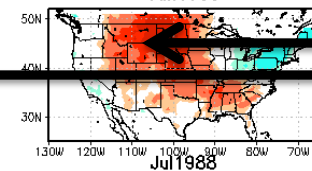
130W 120W 110W 100W 90W 80W 70W



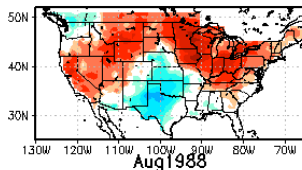
Jun



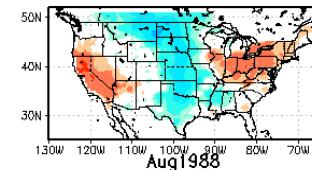
130W 120W 110W 100W 90W 80W 70W



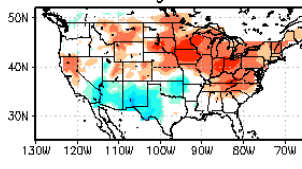
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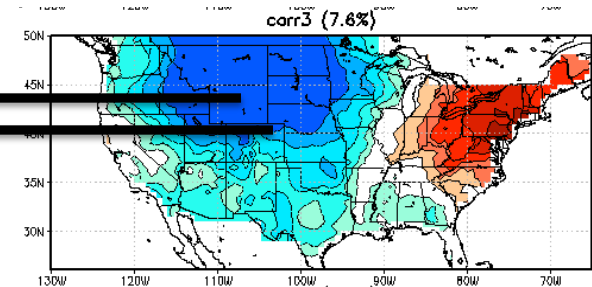
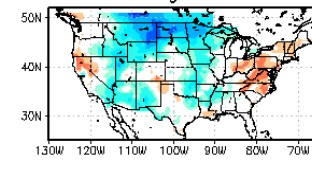
130W 120W 110W 100W 90W 80W 70W



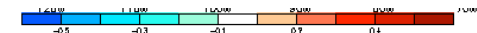
Aug



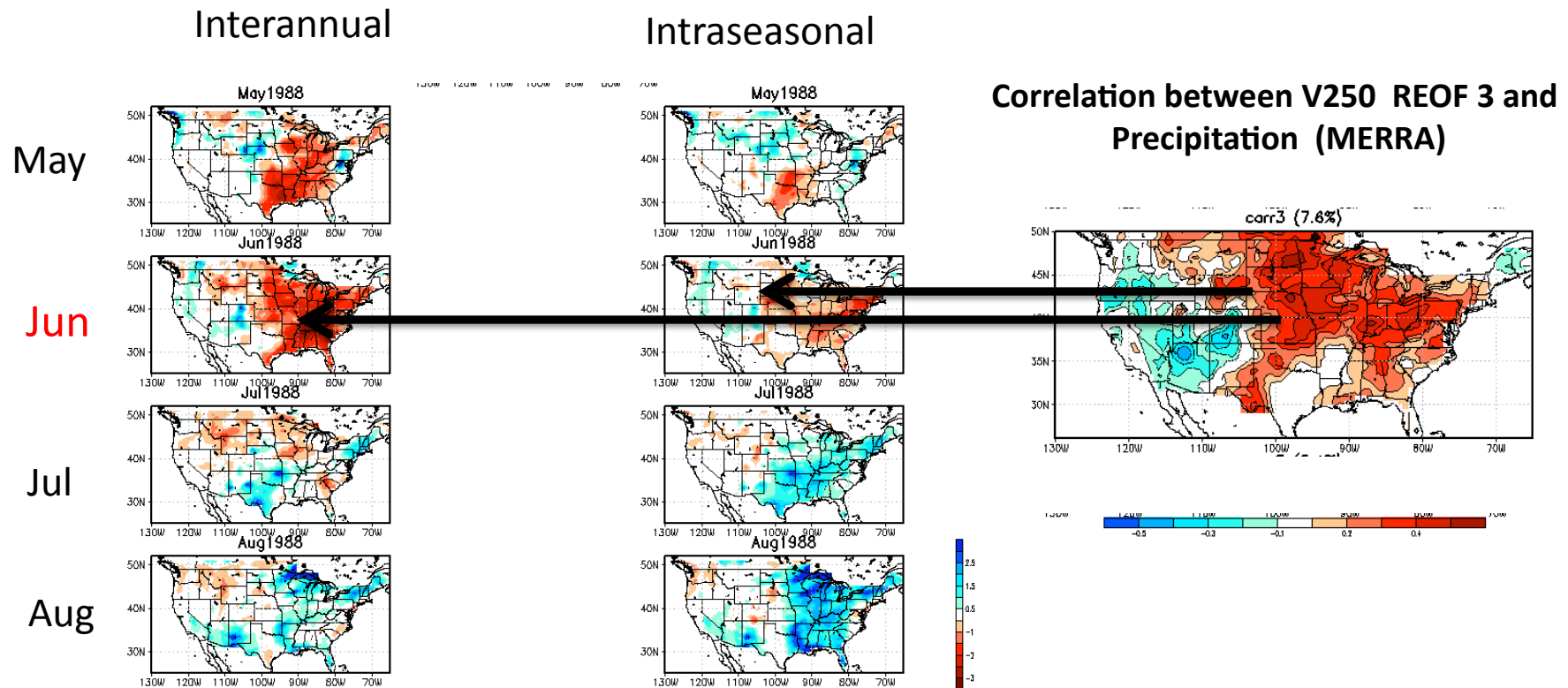
130W 120W 110W 100W 90W 80W 70W



corr3 (7.6%)

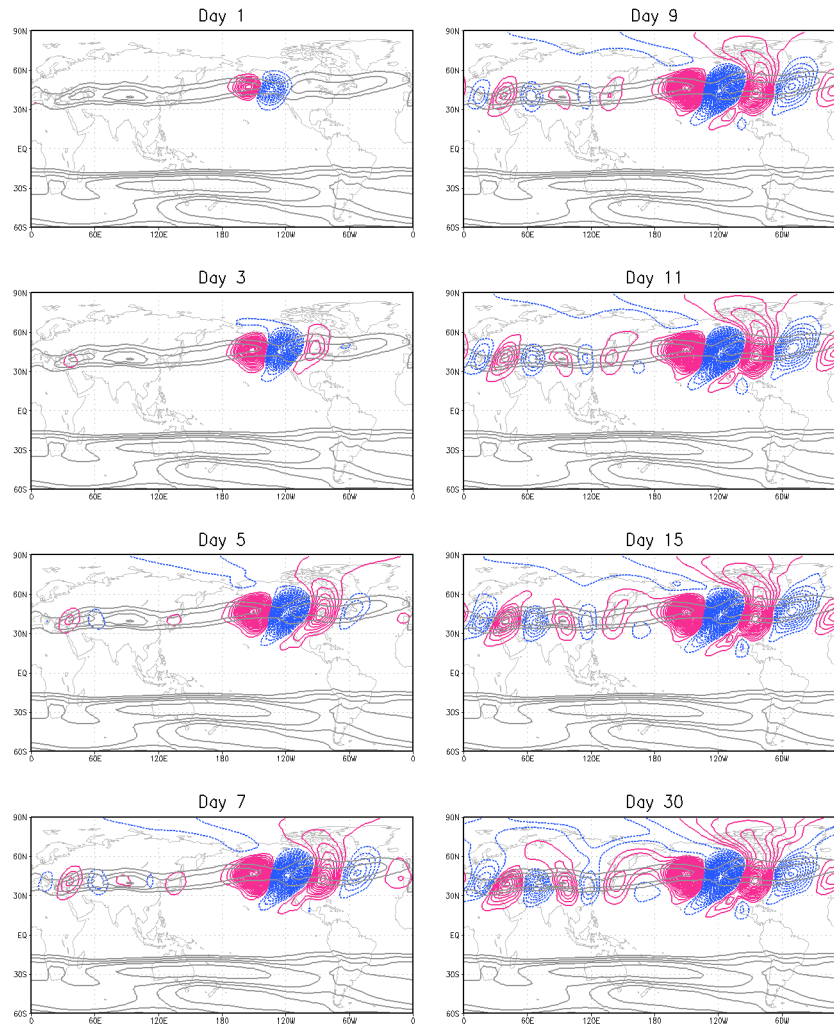


1988 Monthly Anomalies (Precipitation)



Preliminary Work With Stationary Wave Model* (SWM)

Evolution of Eddy V-wind $\sigma=.257$
(Heating at 210E, 45N)



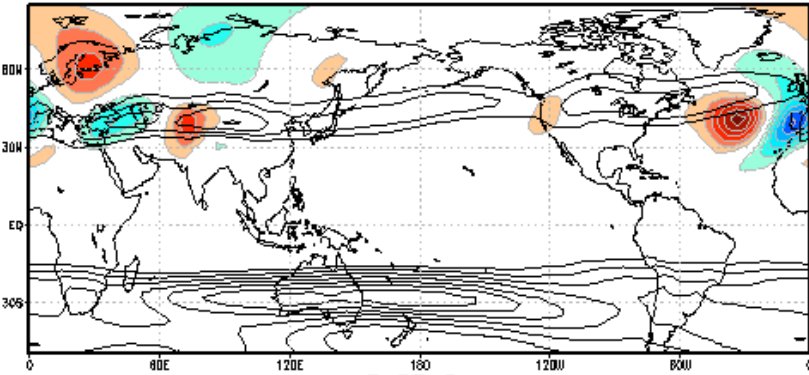
MERRA JJA Base State

*Ting and Yu 1998

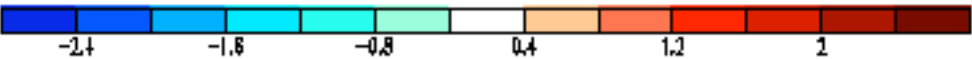
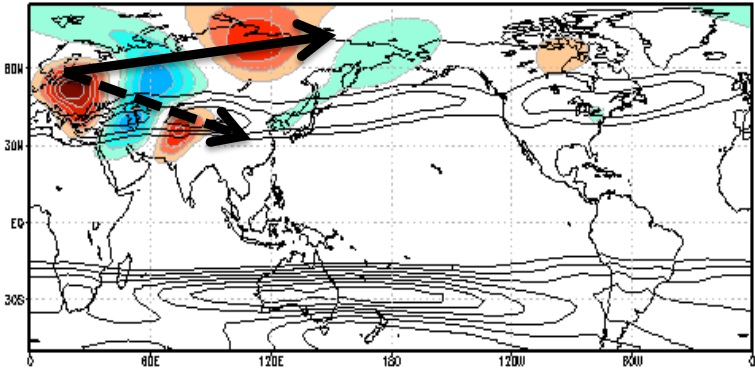
SWM - Examples of responses to idealized forcing in N. Atlantic and Europe

Eddy V-wind $\sigma=.257$

SWM Response to Idealized Heat Source at 330E, 40N



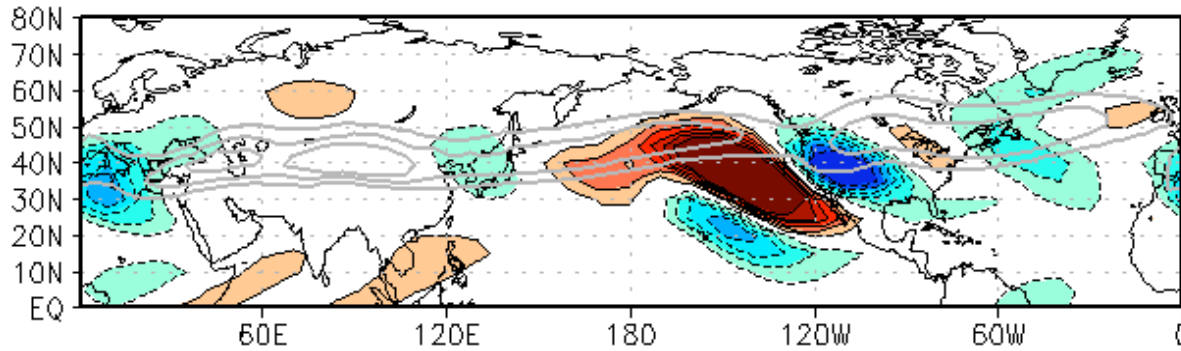
SWM Response to Idealized Heat Source at 20E, 50N



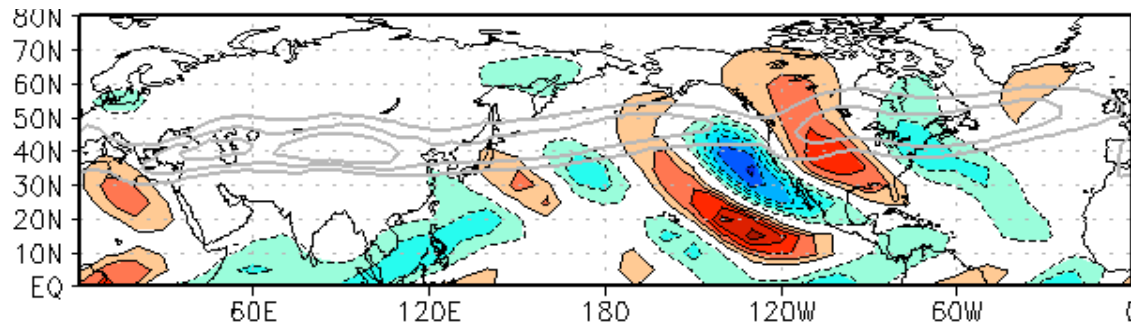
Resembles REOF 1

SWM -Example of optimal forcing patterns for REOF3

“Optimal” Heat Source Pattern

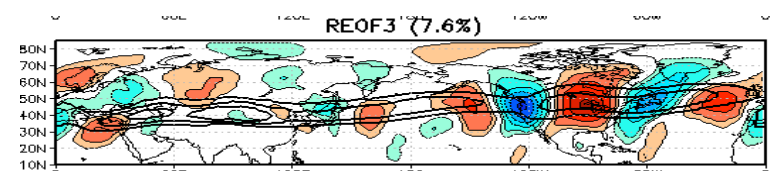


“Optimal” Vorticity Source Pattern



Optimal pattern is computed by calculating the responses to forcings located at every 5° lat and 10°lon and taking the inner product between the response and REOF3 and plotting that at each location

REOF 3

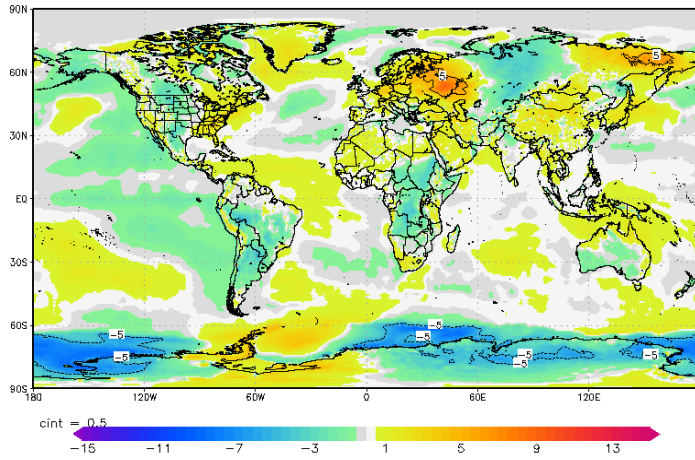


Conclusions

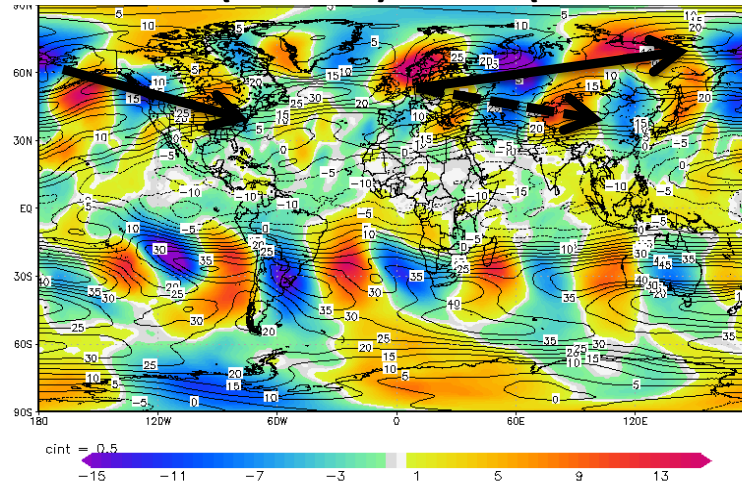
- Stationary Rossby Waves appear to play a significant role in modulating summer precipitation and Tsfc in northern middle latitudes on monthly time scales
- At times, they appear to play a dominant role in the development of regional extremes (e.g., 2003 European heat wave, 1988 US drought)
- What about 2010? During July a Rossby wave seems to play a key role in the extremes in Russia (drought), Pakistan (floods), US Great Plains (very wet) and east coast (very hot) – see next figure.

July 2010 Anomalies

T2m

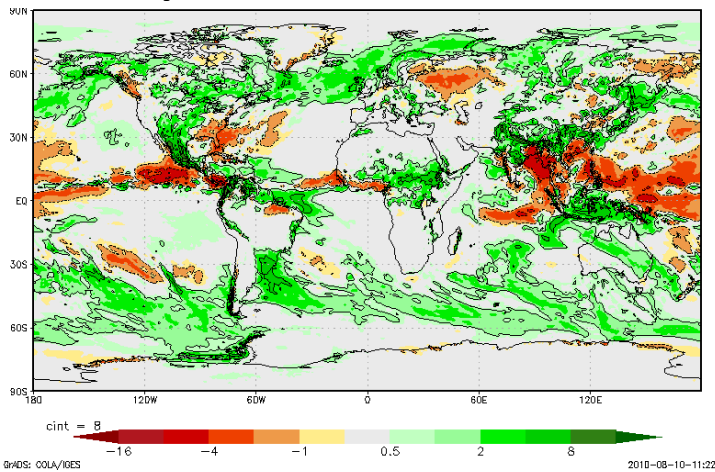


250mb v (shaded) and u (contoured)

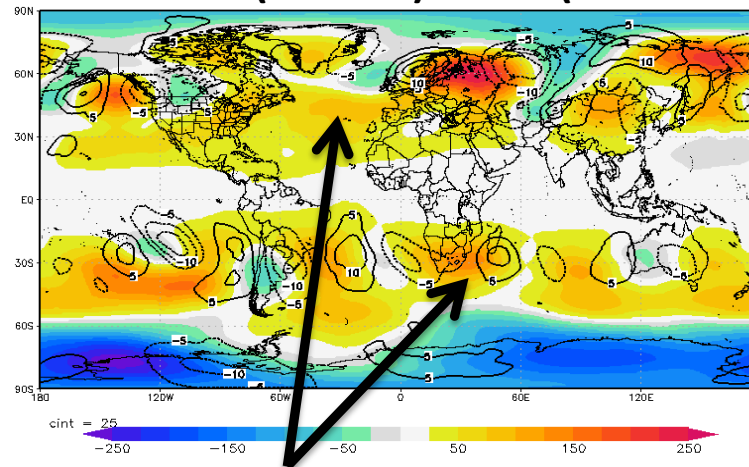


Extremes in Russia, Pakistan, US GP and east coast are all connected by Rossby Wave?

Precip



250mb z (shaded) and v (contoured)



Drought signature: zonally symmetric positive height anomalies in both hemispheres— forced by cold Pacific and warm Atlantic, but appears to have little impact on the July rainfall except regionally in concert with the Rossby Wave

Overall Conclusions

- We have made major strides in the last decade or so
 - Key role of SST (especially tropical Pacific) at interannual/ longer time scales
 - Importance of seasonality in the response to SST
 - Changes in planetary waves/storm tracks in winter and spring
 - Changes in Walker circulation in summer and fall
 - Importance of transition seasons on decadal time scales
- Key Uncertainties
 - Role of Atlantic (and Indian Ocean) SST less clear
 - Model uncertainties (land impacts, Walker circulation response, resolution issues – regional phenomena, weather, basic seasonal cycle, ocean variability in coupled models)
 - Impact of climate change versus decadal variability, aerosols, vegetation, etc
- Issues limiting our ability to predict/monitor
 - SST provide only a relatively weak “background forcing” with a substantial level of noise (stochastic component is large – but presumably depends on time scale)
 - Especially need to address the fact that there is a substantial subseasonal component associated with Rossby waves likely not directly forced by SST that is responsible for short-term modulation of drought/heat waves and other extremes