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2 **Appendix to ‘Probabilistic Precipitation Forecast Postprocessing Using Quantile**

3 **Mapping and Rank-weighted Best-Member Dressing:**

4 **Description of the CSGD method.**

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6 Here we describe the modifications of the Censored, Shifted Gamma Distribution
7 method by Scheuerer and Hamill (2015) made in order to address the particular challenges of
8 this study.

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10 The original CSGD approach started with quantile mapping the forecasts, enlarging the
11 ensemble by including forecasts at nearby grid points, and calculating a number of ensemble
12 statistics. We perform the same three steps here, but in doing so we proceed as described in
13 Section 3b of this paper and *not* as proposed by SH15. Some of the ensemble statistics
14 considered here are also slightly different: we still use the ensemble mean \bar{x}^f , and the fraction
15 of non-zero ensemble members $F \hat{N} Z$, but as a measure of ensemble spread we use the
16 standard deviation $\sigma(\tilde{x}^f)$ instead of the mean absolute difference. We do not use precipitable
17 water as a predictor here.

18 The second step in the procedure described by SH15 is to fit climatological CSGD
19 parameters - separately for each month, each lead time, and each grid point - to the analyzed
20 precipitation amounts used for training. Here, we do this using analysis data from 2002 to 2017
21 and a slightly different model fitting approach. As in SH15, for each month the training data is
22 composed of the 45 days before and after the 15th of each month. From these data, we

23 calculate the climatological mean \bar{y}_{cl} and the climatological fraction of zero (here < 0.254 mm)
 24 precipitation analyses $\hat{F}Z_{cl}$. We choose the shape parameter k_{cl} of the climatological CSGD
 25 such that the continuous ranked probability score (CRPS) over the training sample is minimized,
 26 while the scale parameter θ_{cl} and the shift parameter δ_{cl} are chosen such that the CSGD
 27 defined by those three parameters has the prescribed climatological mean \bar{y}_{cl} and
 28 climatological fraction zero $\hat{F}Z_{cl}$. Specifically, denote by F_k the cumulative distribution
 29 function (CDF) and by f_k the probability density function of a gamma distribution function with
 30 shape parameter k . Let F_k^{-1} be the inverse CDF and define $q_0 := F_k^{-1}(\hat{F}Z_{cl})$. For a
 31 given k , the parameters θ_{cl} and δ_{cl} can then be calculated via

$$\theta_{cl} = \frac{\bar{y}_{cl}}{k(1 - \hat{F}Z_{cl} + f_{k+1}(q_0)) - q_0(1 - \hat{F}Z_{cl})}, \quad \delta_{cl} = -\theta_{cl} q_0$$

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 35 The climatological shape parameter k_{cl} is then found via CRPS minimization as described in
 36 SH15. In contrast to their original suggestion, minimization is now performed over a 1-
 37 dimensional (instead of 3-dimensional) parameter space, and is therefore computationally more
 38 efficient. Moreover, the three quantities k_{cl} , \bar{y}_{cl} and $\hat{F}Z_{cl}$ have an intuitive interpretation and
 39 can be assumed to have a smooth annual cycle. We can therefore linearly interpolate them from
 40 the 15th of each month to every single day of the year, and calculate the climatological CSGD
 41 parameters μ_{cl} , σ_{cl} and δ_{cl} from eqs. (1), (2) in SH15 from the interpolated values of k_{cl} , \bar{y}_{cl} ,
 42 and $\hat{F}Z_{cl}$.

43 The final step in SH15 is to link the CSGD parameters of the calibrated forecast
 44 distribution at grid point s to the ensemble statistics at s defined in the first step. In the present
 45 setup, we use the following regression equations

$$46 \mu_s = \frac{\mu_{cl,s}}{\alpha_1} \log 1p \left[\text{expm1}(\alpha_1) \left(\exp(-\rho_s/\alpha_2) + \alpha_3 F\hat{N}Z_s + \alpha_4 \rho_s \frac{\tilde{x}_s^f}{\tilde{x}_{cl,s}^f} \right) \right]$$

$$47 \sigma_s = \alpha_5 \sigma_{cl,s} \sqrt{1 - \rho_s^2} \sqrt{\frac{\mu_s}{\mu_{cl,s}}} + \alpha_6 \sigma(\tilde{x}_s^f),$$

$$48 \delta_s = \delta_{cl,s},$$

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 50 where $\log 1p(x) = \log(1 + x)$, $\text{expm1}(x) = \exp(x) - 1$, and $\tilde{x}_{cl,s}^f$ is the
 51 climatological average of the ensemble mean \tilde{x}_s^f of the quantile-mapped forecasts. Due to the
 52 quantile mapping, we can assume $\tilde{x}_{cl,s}^f = \bar{y}_{cl,s}$, where $\bar{y}_{cl,s}$ is the climatological mean of the
 53 analyzed precipitation at s . The regression equations above differ from those used in SH15 in a
 54 number of ways. First, there are only six regression parameters since there is no precipitable
 55 water predictor and since the heteroscedasticity parameter was fixed to 0.5 (a simplification
 56 suggested by SH15). Second, due to the limited training sample size and the challenges that
 57 come with the estimation of a rather complex model, the regression parameters $\alpha_1, \dots, \alpha_6$
 58 are assumed constant across the entire domain, and estimated in a single CRPS minimization.
 59 As explained in SH15, one of the motivations for incorporating the climatological CSGD
 60 parameters $\mu_{cl,s}$ and $\sigma_{cl,s}$ into the regression equations is that this removes some of the local
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62 characteristics. One characteristic that is not captured that way is the local forecast skill of the
63 NWP model, which can vary substantially across the domain. By assuming $\alpha_1, \dots, \alpha_6$
64 constant in space, these parameters cannot account for spatially varying skill either, and this
65 motivated the inclusion of a spatially varying skill parameter ρ_s . This parameter is defined as the
66 correlation between the mean of square-root transformed, quantile-mapped ensemble forecasts
67 and square-root transformed analyzed precipitation amounts. The calculation is performed with
68 the same training sample used for estimating the regression parameters, but each analysis grid
69 point was supplemented by the best 19 supplemental locations found as described in Section
70 3a. The particular way of including ρ_s in the regression equations above is motivated by
71 standard regression theory. If two variables have correlation ρ , the slope parameter for
72 regressing one of them on the other is proportional to ρ , and the unexplained variance is
73 proportional $1 - \rho^2$. Relating the intercept parameter to ρ_s is more challenging in the present
74 context; we chose the expression $\exp(-\rho_s/\alpha_2)$ since it ensures that the intercept is always
75 positive, is equal to 1 if there is zero skill, and tends to zero as skill increases. With these three
76 changes, we hope to strike a balance between parsimonious parameterization - which is
77 necessary in a situation with limited training data - and sufficient flexibility to address local
78 characteristics of the different analysis grid points within the CONUS.

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